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A Seismic/Acoustic Study of Low Altitude Air Blasts at Ft Devens, Massachusetts

JANET C. JOHNSTON



9 March 1990



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JAME'S LEWKOWICZ

Earth Sciences Division

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Division Director

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COSATI CODES 18. SUBJECT TERMS (Continue on reverse if necessary and identify by block number) Air Blasts, (160) Seismic/Acoustic → Aerial Blasts. SUB-GROUP FIELD **GROUP** Aerial Explosions Seismic Detection

9. ABSTRACT (Continue on reverse if necessary and identify by block number)

On 5 May 1986, 3 packages of TNT were detonated 100 ft in the air above the Turner Drop Zone at Ft. Devens, Massachusetts. The aerial explosions were recorded by seismic and acoustic sensors over 2 linear arrays that covered the distance range from 0 to 1000 meters. One array was installed over flat topography, while the other was installed across a hill with 80 ft of relief. The goal of the experiment was to examine the response of the ground and the air to a unit impulse. Although recording of the higher frequencies generated by the blasts was limited by the sampling rate capability of the equipment (100 and 200 samples/sec) it was clear that the hill produced a sheltering effect. Considerable transverse motion appeared in the records. Reduced time plots revealed a propagation speed of 0.35 km/sec corresponding primarily to an airwave induced seismic signal, although some seismic precursors are evident. The ratio of peak seismic velocity to pressure ranged from 10⁻⁴ cm/sec/Pa for the line that transversed the hill to 10⁻⁹ cm/sec/Pa for the flat array. The experiment should be repeated in the same area during different seasonal coverings (snow and ice) and extended to areas with more complex topography and different vegetation.

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UNCLASSIFIED SECURITY CLASSIFICATION OF THIS PAGE In future experiments the position of the air blasts should be varied to facilitate discrimination of paths to the receivers. The results could be extrapolated to an extended, continuous source such as a low flying aircraft or cruise missile.

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Preface

The author thanks the many people whose cooperation made this experiment possible. These include Ken Dallas, who coordinated balloon support, Bob Reinke and Al Leverett from the Weapons Laboratory, all Solid Earth Geophysics Branch personnel who helped install the equipment and students and other last minute assistants; Katharine Kadinsky-Cade, Joe Sabeni, Janet Oliveira-Jones and Dave Shorter. Also, locations were surveyed by Ted Writanen (formerly of the Geodesy and Gravity Branch) and Lee Stevens of GL did a fine job of photography as well as digging holes.

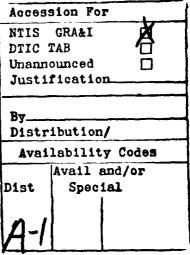
A special thanks goes to SMSgt Roger Sands (Geodesy and Gravity Branch) without whose assistance the experiment would not have been possible.

The author benefited from many helpful discussions of seismo-acoustic interaction with James C. Battis of the Solid Earth Geophysics Branch (GL).

Thank you also to John Cipar (Solid Earth Geophysics Branch, GL) and James C. Battis for reviewing this manuscrip.

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A Seismic/Acoustic Study of Low Altitude Air Blasts at Ft Devens, Massachusetts

1. INTRODUCTION

On 5 May 1986, three packages of dynamite, two 5-pound bundles and on 10-pound bundle, were detonated in the Turner Drop Zone area on Fort Devens Army Base in eastern Massachusetts. The packages were suspended approximately 100 feet in the air by balloon. The ground motions and airwaves resulting from the three separate shots were recorded by a digital array consisting of 10 seismic and 9 acoustic stations. A crew of 25 scientists, engineers, and technicians from the Earth Sciences Division, Aerospace Instrumentation Division (balloon support) of the Geophysics Laboratory, and the Civil Engineering Research Division of the Weapons Laboratory assembled and installed the equipment in one day at the Turner Drop Zone.

The goal of this experiment was to examine the response of the ground to what may be considered essentially a unit impulse. The results may be then extrapolated to an extended, continuous source such as a low flying aircraft or a cruise missile. The experiment was designed to be repeated in the same area during different seasonal coverings (snow and ice) and extended to areas of more complex topography, terrain and differing vegetation.

2. GENERAL GEOLOGIC SETTING AND TOPOGRAPHIC INFORMATION

Fort Devens lies between the Essex Fault and the Pine Hill Fault in northeastern Massachusetts (Figure 1). The Essex Fault is a regionally developed, east-northeast trending dislocation that separates the Merrimac Group on the north and west from a possibly Precambrian and generally much more highly metamorphosed terrain on the south and east. The Pine Hill Fault generally marks the boundary between the Merrimac Group and Worcester Phyllite. The Merrimac Group consists of layered schists and gneisses of probable Ordovician-Silurian age. Also, the Ayer granite (Ordovician) has been mapped in the quadrangle. As with all of central and eastern Massachusetts, the structural characteristics result from a long and complex tectonic history. The drop zone is a relatively flat open field with a hill of 80-ft relief in its northeast corner. Both the field and the hill lack significant vegetation. There were a few trees and a covering of grass. The soil was rocky and some large boulders were apparent on the north side of the hill. The edge of a wooded zone ran north-south about 75 meters from ground zero. In the immediate area of the experiment, (see topographic map, Figures 2 & 3)³, it was possible to lay out a relatively flat east-west profile (X-line) in contrast to a north-south profile (Y-line) that stretched over a hill with a relief of 80-ft (25 m). These profiles and station locations, as well as the shot position in the air, were folded onto two dimensions in Figure 3.

3. SEISMO-ACOUSTIC ARRAY

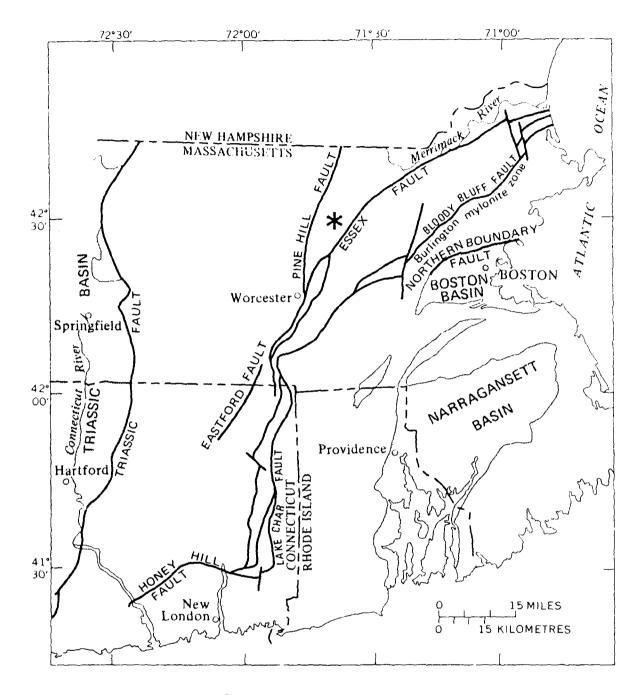
The three-component seismic ground motions and air pressures from the blasts were recorded digitally on an array consisting of two approximately radial profiles (Figure 3). One profile ran approximately north of the shots (Y-line), the other approximately west (X-line) of the shots. The shot point is referred to as "ground zero" for the remainder of this report. The X-line consisted of six 3-component seismic and pressure stations, and the Y-line consisted of four 3-component seismic stations and three pressure stations (at station locations: Y1, Y2, Y3).

Timing was accomplished by setting all the recorder clocks to a master clock before the first shot. Recorder clock offsets were measured with respect to a single recorder after the experiment (Appendix A, Table A-1). Geospace HS-10 seismometers with a natural period at 1 Hz and Sprengnether S-6000 seismometers with a natural period of 2 Hz were connected to the Terra Technology DCS-302 seismic recorders. The seismometers were buried in shallow holes to minimize wind noise. Pressure was recorded by Model MLR pressure transducers (± 0.5 PSI full scale and ± 1.5 PSI full scale). For some of the stations, in-ground calibration tests were run as teardown time permitted. Recorder/seismometer pairs were calibrated in the laboratory, as were all the pressure

^{1.} Castle, R.O., Dixon, M.R., Grew, E.S., Griscom, A., and Zietz, I. (1976) Structural Dislocations in Eastern Massachusetts, Geological Survey Bulletin 1410, U.S. Government Printing Office, Washington, DC.

^{2.} Van Diver, B.B. (1987) Roadside Geology of Vermont and New Hampshire, Mountain Press Publishing Company, Missoula, pp.1-29.

^{3.} United States Army Forces Command, 584th Engineer Company (Cartographic), 30th Engineer Battalion (Topographic) (Army), Fort Belvoir, VA 22060, (Feb. 1986) Fort Devens Special Map, edition 2-DMATC, Series V8145.



* Location of Fort Devens, Turner Drop Zone

Figure 1. Structural Dislocations in Eastern Massachusetts (Modified from Castle et al. 1976)

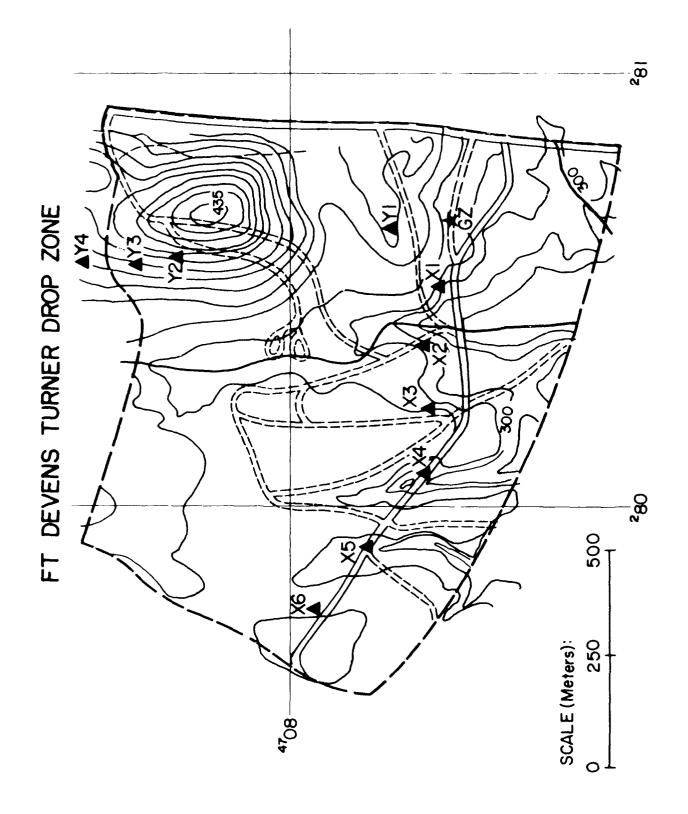


Figure 2. Site Map with Topographic Contours

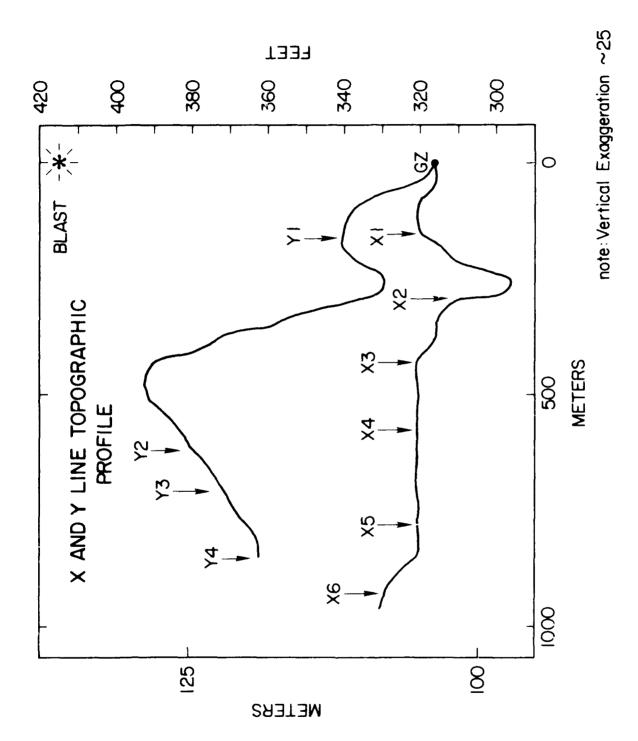


Figure 3. X- and Y-Line

transducers (Appendix A, Table A-2). The Terra Technology DCS-302 recorders used cassette tapes to record the signals. For the majority of the stations the recorders were triggered remotely using cables. Some stations were set on trigger manually a few seconds before shot times.

The shots were recorded at either 200 samples per second or 100 samples per second. Table A-3, in Appendix A contains information on sampling rates for each station and Appendix B contains additional information on the instruments for the X- and Y-lines. The response of the recorder's amplifier is flat from DC to the anti-aliasing filter cut-off frequency of 30 Hz or 70 Hz. Above that cut-off frequency, the response rolls off at 6 dB per octave using a 6-pole Butterworth filter. The recorders that were attached to seismometers were set at automatic gain ranging; the recorders attached to the pressure sensors were set to fixed gain.

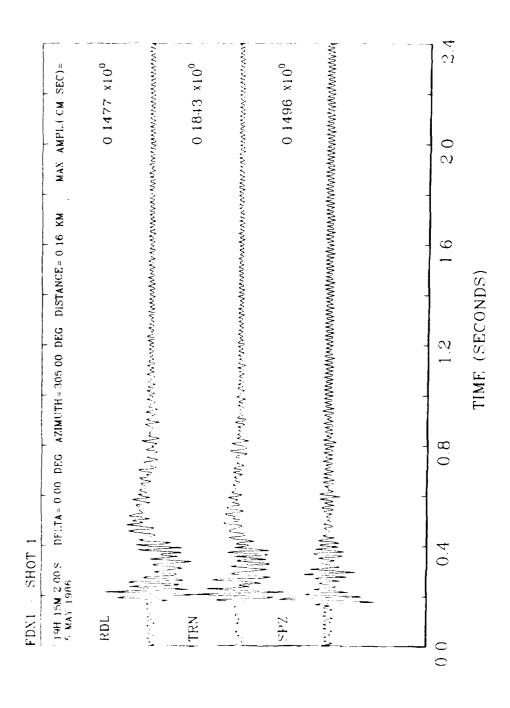
4. ANALYSIS AND INTERPRETATION

Table 1 lists distances to ground zero and azimuths for each station. The seismogram recordings

Station Distance to Map Elev. Bearing Back Map Symbol G.Z. (Meters) (Meters) from G.Z. Azimuth (Magnetic) 105.156 00° Y1 FDY1 147.218 180° Y2 FDY2 643.433 123.444 00° 180° 00° 180° **Y3** FDY3 744.626 117.348 15° 192° Y4 FDY4 850.697 114.300 G.Z. Porta Corder 0 97.536 94.488 305° 125° X1FDX1 160.325 X2 309.677 91.44 305° 125° FDX2 X3 FDX3 445.008 96.012 280° 100° X4 FDX4 593.750 94.488 295° 115° X5 742.493 97.536 300° 120° FDX5 X6 FDX 6 891.235 97.536 305° 125°

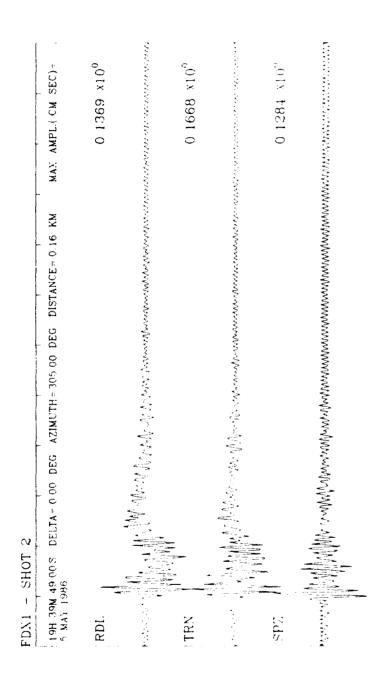
Table 1. Station Location Data

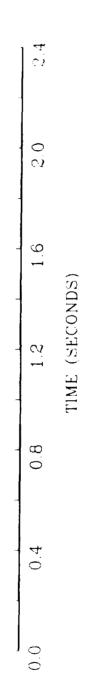
were rotated into radial and transverse components. Figures 4 and 5 display all the seismic waveforms from the X- and Y-line respectively in order of station location. Amplitude spectra were computed for vertical components and can be found in Appendix C. In Figures 4 and 5, the numbers positioned on the right sides of the plots give the maximum amplitude of the signal in cm/sec. The high frequency background on station FDX1 and FDY2 for all shots resulted from a malfunction internal to the recorder. The very long period motion at station FDX1 radial component is probably a ground roll phenomenon. FDX1 was one of the closest stations to the shots at 160 m. However, this



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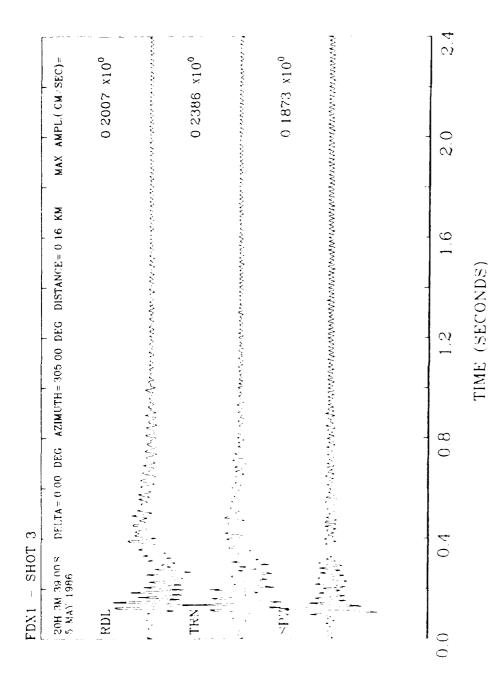
Figure 4. Seismograms - X-Line





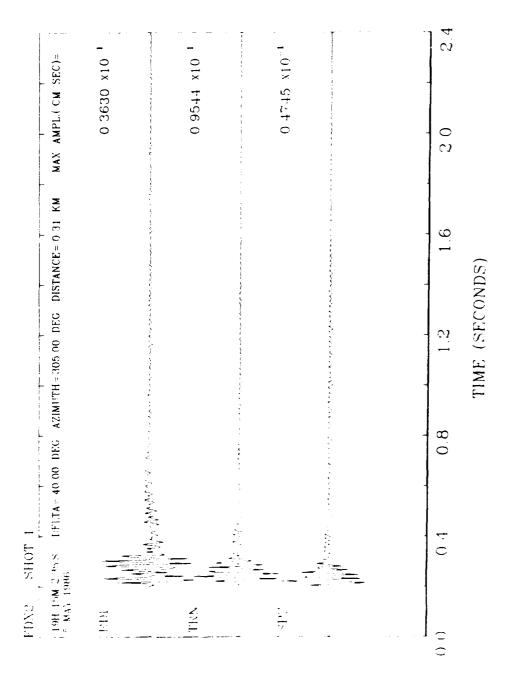
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Figure 4. Seismograms - X-Line (Cont'd)



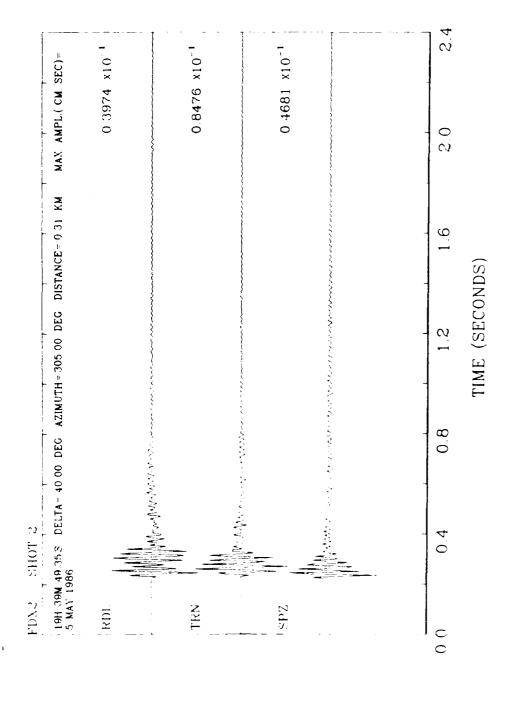
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Figure 4. Seismograms - X-Line (Cont'd)



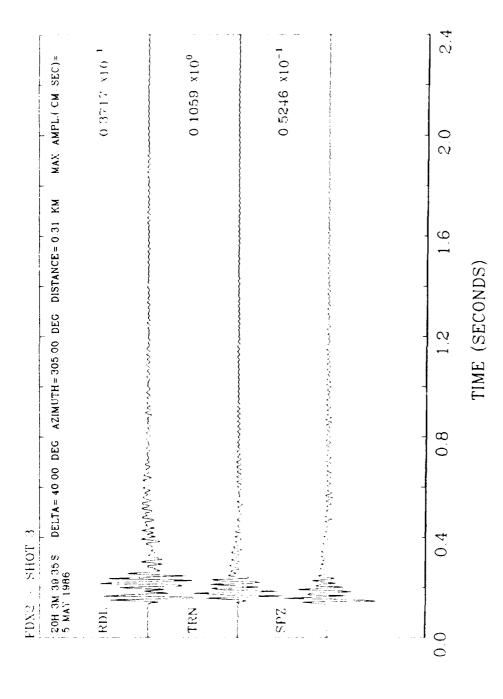
11-JUL-88 14-02-11

Figure 4. Seismograms - X-Line (Cont'd)



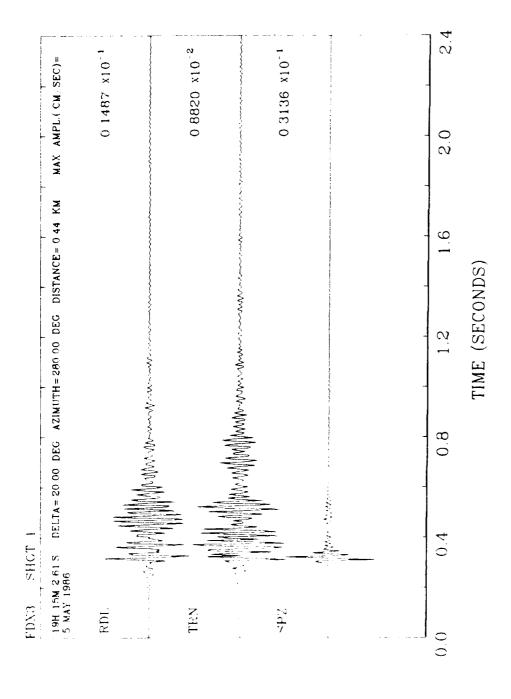
11-JUL-88 14-03-26

Figure 4. Seismograms - X-Line (Cont'd)



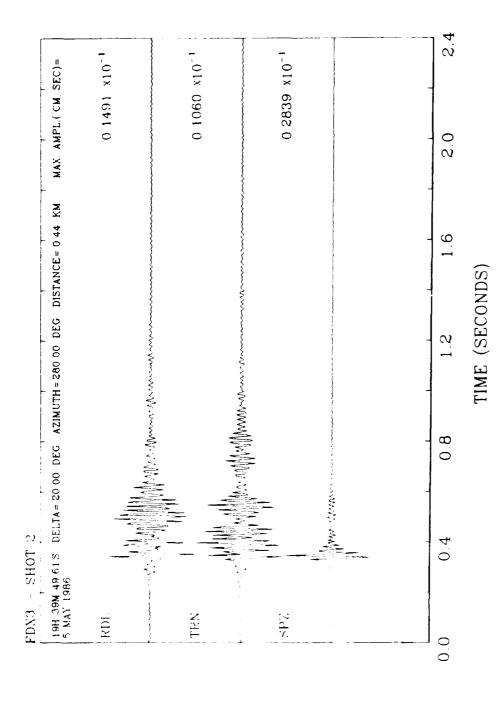
11-JUL-88 14-04-29

Figure 4. Seismograms - X-Line (Cont'd)



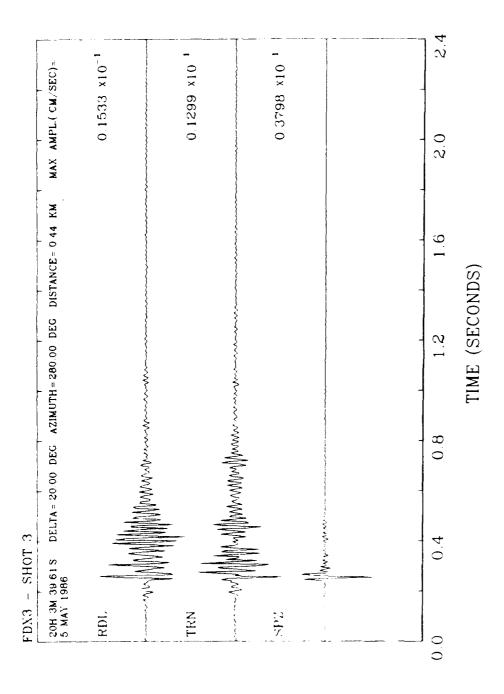
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Figure 4. Seismograms - X-Line (Cont'd)



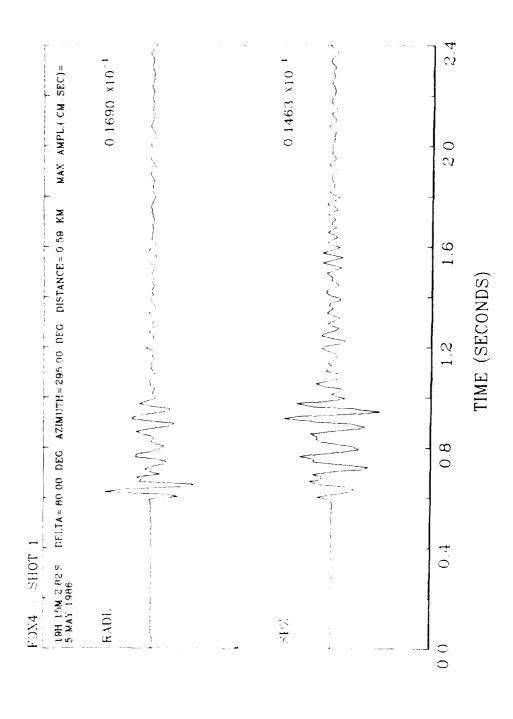
11-JUL-88 14-07-13

Figure 4. Seismograms - X-Line (Cont'd)



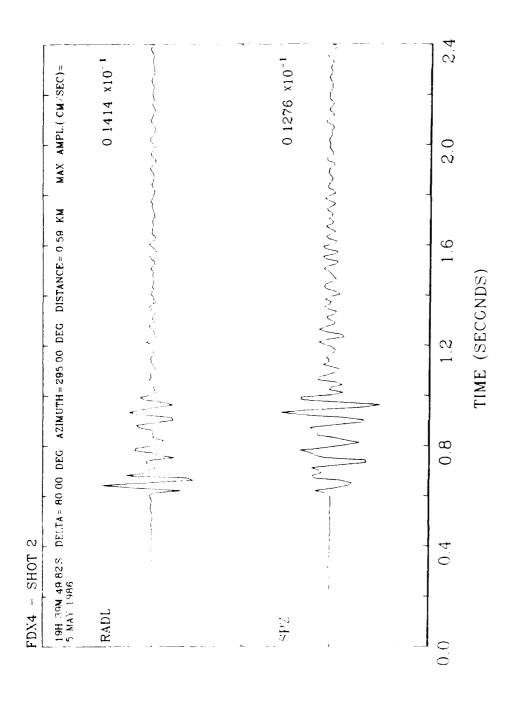
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Figure 4. Seismograms - X-Line (Cont'd)



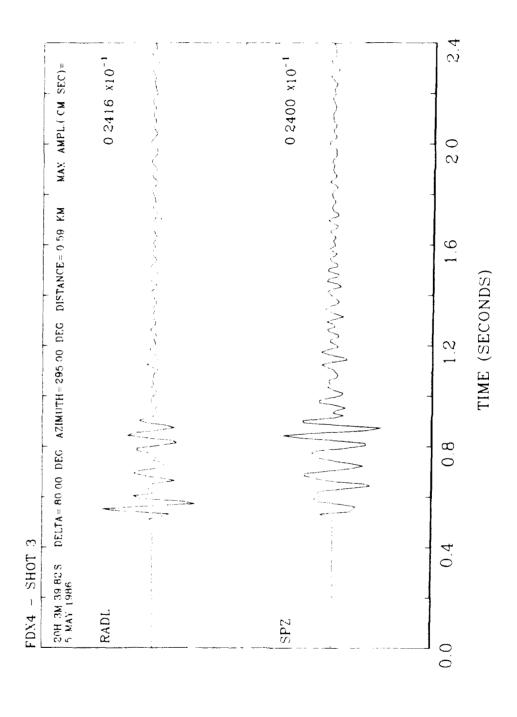
11-JUL-88 14-09-34

Figure 4. Seismograms - X-Line (Cont'd)



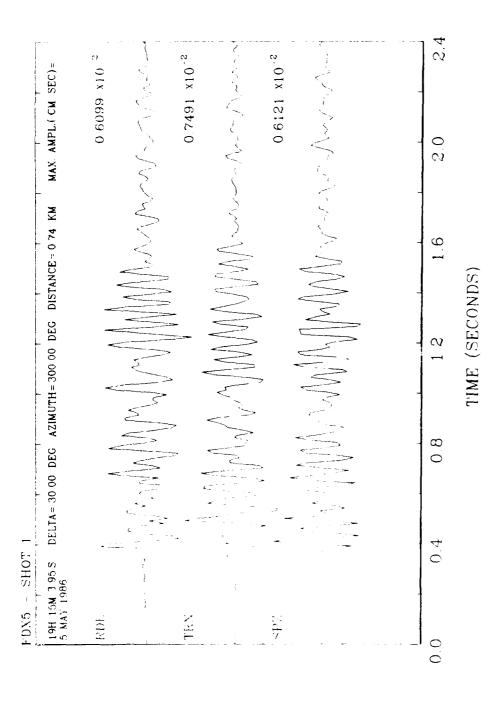
11-JUL-88 14-10-45

Figure 4. Seismograms - X-Line (Cont'd)



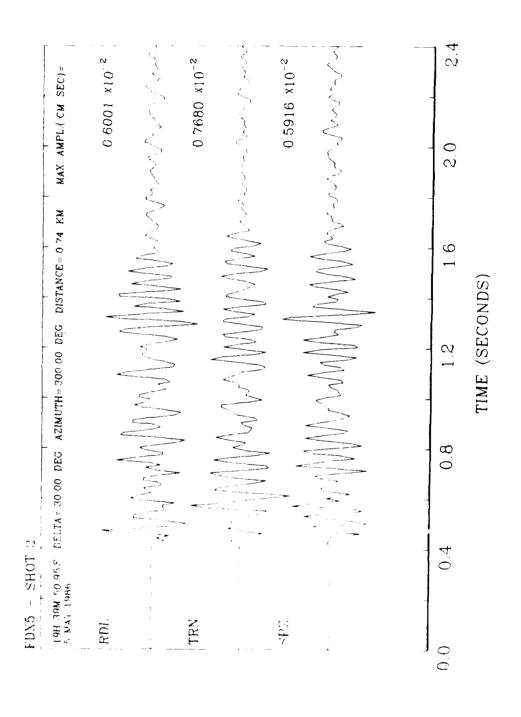
11-JUL-88 14-11-50

Figure 4. Seismograms - X-Line (Cont'd)



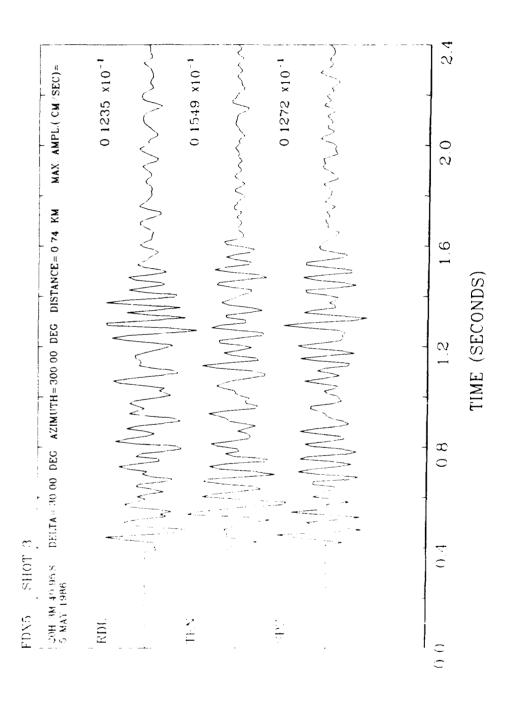
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Figure 4. Seismograms - X-Line (Cont'd)



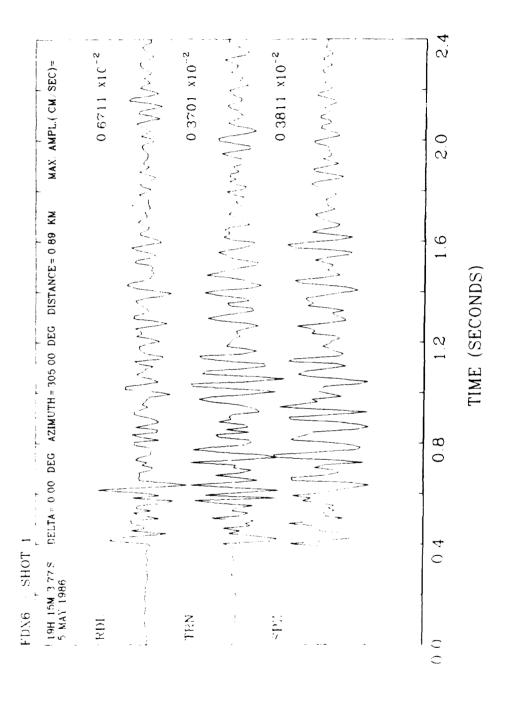
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Figure 4. Selsmograms - X-Line (Cont'd)



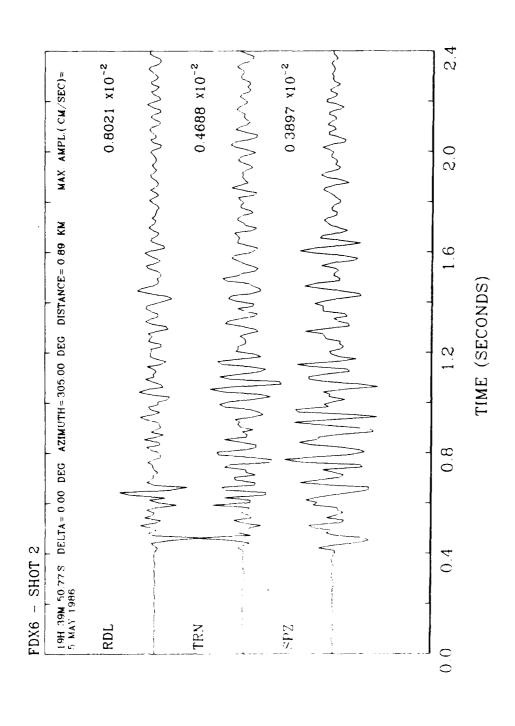
11-JUL-88 14-16-21

Figure 4. Seismograms - X-Line (Cont'd)



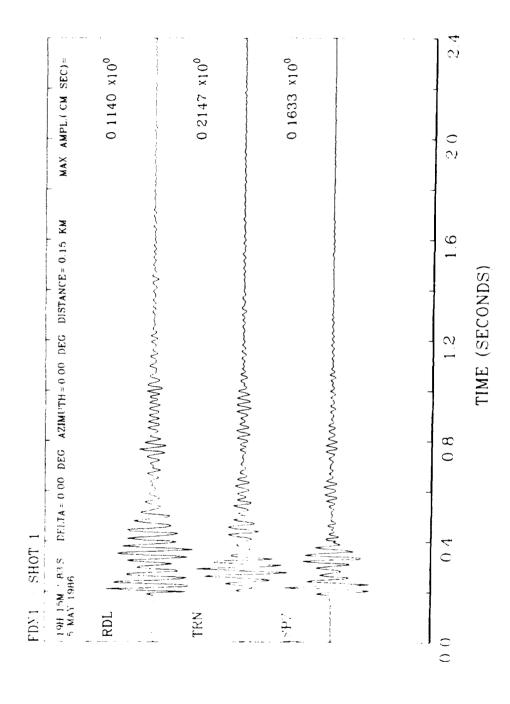
11-JUL-88 16-46-54

Figure 4. Seismograms - X-Line (Cont'd)



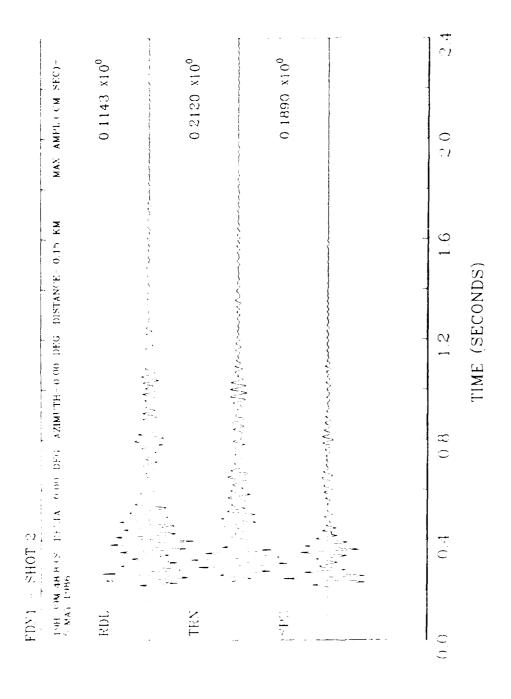
11-JUL-88 14-18-12

Figure 4. Seismograms - X-Line (Cont'd)



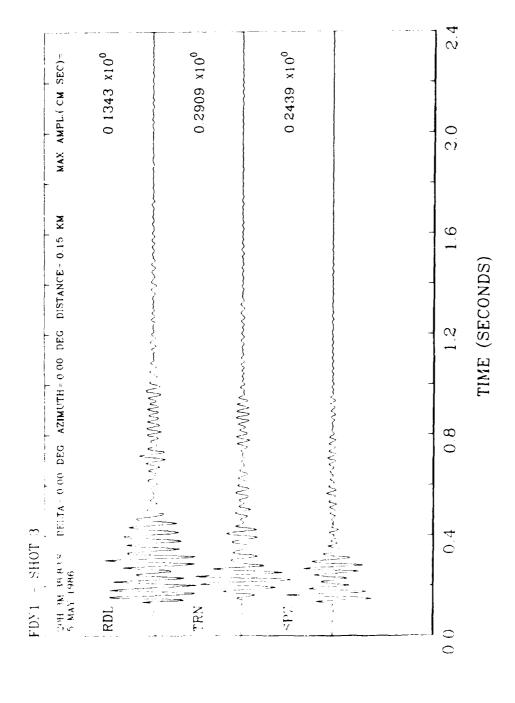
7 JUL 88 15 10 55

Figure 5. Seismograms - Y-Line



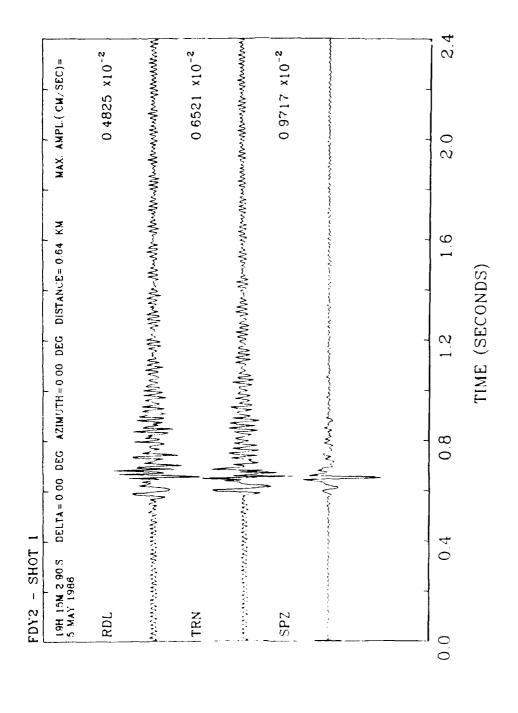
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Figure 5. Seismograms - Y-Line (Cont'd)



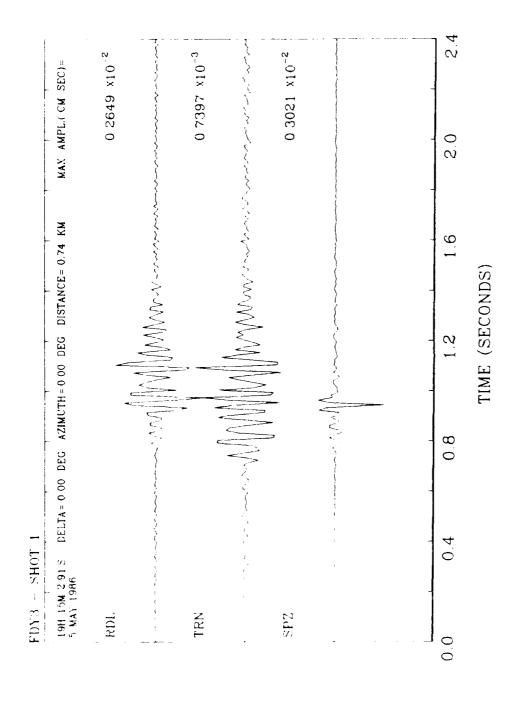
7-JUL-88 15-18-54

Figure 5. Seismograms - Y-Line (Cont'd)



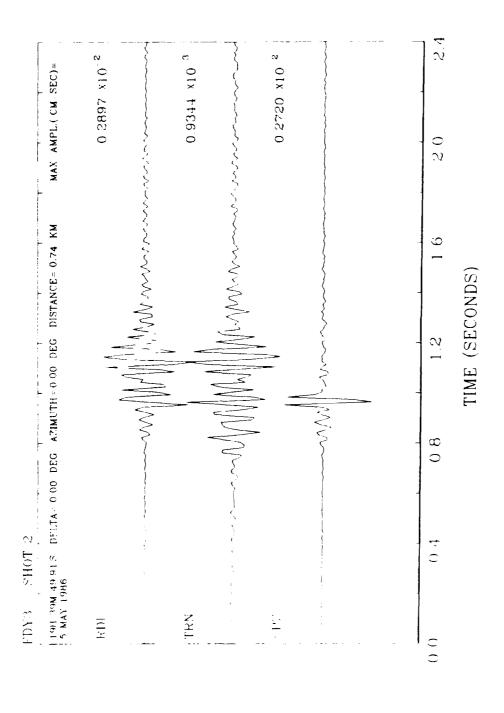
7-JUL-88 16-36-12

Figure 5. Seismograms - Y-Line (Cont'd)



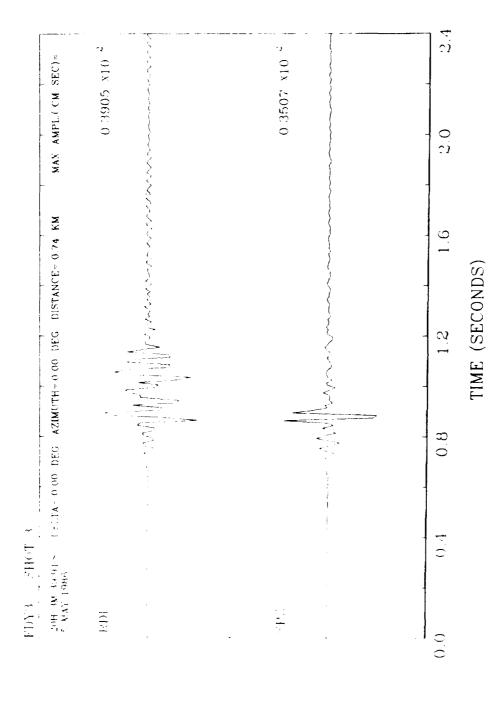
7-JUL-88 16-38-57

Figure 5. Seismograms - Y-Line (Cont'd)



7 · JUL ·88 16--40-00

Figure 5. Seismograms - Y-Line (Cont'd)



7-JUL-88 16-41-04

Figure 5. Seismograms - Y-Line (Cont'd)

effect is not seen at the closest Y-line station, FDY1. Generally, the radial component was of marginally larger amplitude than the transverse and the vertical. Explosive sources at these distances usually generate considerable radial and vertical motion. Maximum amplitudes for shots 1 and 3 (all 3 components) are plotted on Figures 6 and 7 for the X- and Y-line.

The amplitudes are scattered, but it can be seen that the X- and Y-lines have similar values on the ground zero side (less than 350 m) of the Y-line hill and the values begin to diverge after that. In fact, at the 750 m range the Y-line amplitudes are less than the X-line amplitudes by a factor of approximately 3 on instruments with similar response characteristics. Attenuation with distance should not be measured from this plot since a change in instrumentation occurs in the middle of the lines.

 $Mederis^4$ fit peak amplitudes of velocity to curve of the form

$$V = aL^bW^c$$
 (1)

where V is the peak ground velocity in inches/per second, L is the distance to a detonation (in feet) and W is the explosive weight in pounds. For the vertical component, the arbitrary constants, a, b, and c are 69.2, -1.37, and 0.59 respectively, with a standard error of 2.10. These values were obtained using in-ground or surface shot data primarily from quarries, but the predictions conform within the error to measurements of the aerial shots used in this report for the flat X-line.

Additionally, ratios of maximum amplitude at the same station for different shots reveal that doubling the amount of explosive increased the observed amplitudes by a factor of 1.1 to 1.5. (The higher factors corresponded to the Y-line stations sheltered by the hill). Nicholls et al⁵ found similar results for scaled distance versus overpressure for a range of shot sizes.

Amplitude spectra for vertical components (shot 1) seismograms that appear in Appendix C show variation of peak energy with frequency. This is a reflection of the different instrument responses. The spikes in the spectra at approximately 50 Hz for FDX1, FDX2, and FDY2 are a result of a malfunction in the recorders (this high frequency noise can be observed on the seismograms). Most of the records exhibit peaks in their seismic spectra between 10 and 25 Hz. An exception is FDX3 shot 1 which peaks at 60 Hz. Completely unexplained is the large dip in the FDY3 spectrum at 22 Hz. A less pronounced dip can be observed on the other Y-line stations. It may be the result of destructive interference between a direct acoustic induced wave at the station and air coupled seismic waves that originate on the blast side of the hill and travel on multiple paths. In a seismic/acoustic study at Vandenberg AFB, Battis⁶ found that for similar sized explosions at a distance of 300 m, the energy peaked at approximately 18 Hz. Peaks at this frequency are evident in some of the spectra, however many of the characteristics remain unexplained at this time. If the experiment were to be repeated, varying the position of the shot, at least in the vertical direction, would help to unravel the complexities of the spectra and seismic waveforms.

^{4.} Medearis, K. (1979) Dynamic Characteristics of Ground Motions Due to Blasting, Bull. Seis. Soc. Am., **69**(No. 2):627-639.

^{5.} Nicholls, M.R., Johnson, C.F., and Duval, W.I. (1971) Blasting Vibrations and Their Effects on Structures, Bulletin 656, US Dept. of the Interior, Bureau of Mines, pp. 65-66.

^{6.} Battis, J.C. (1985) Vibro-Acoustic Forecasts for STS Launches at V23, Vandenberg AFB: Results Summary and the Payload Preparation Room, AFGL-TR-85-0013, ADA162192.

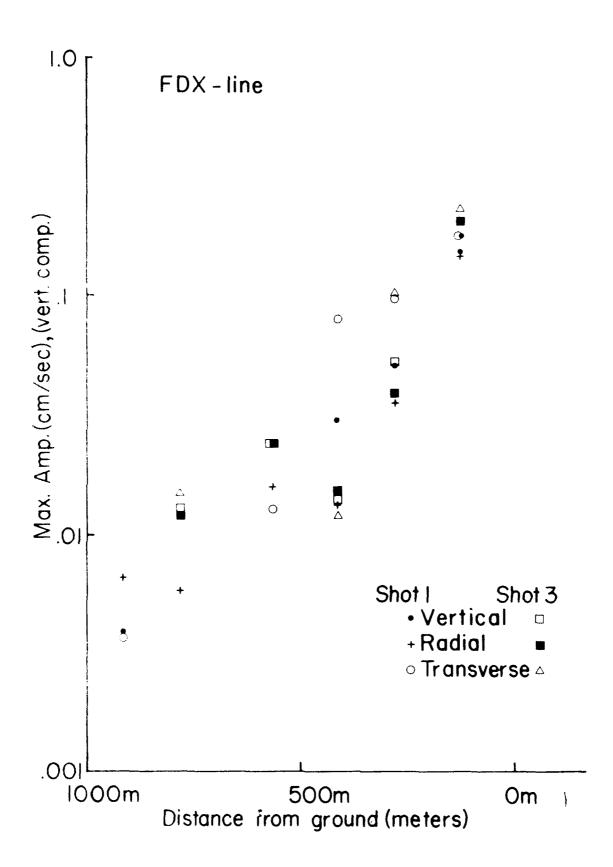


Figure 6 - Maximum Amplitude vs Distance - X-Line

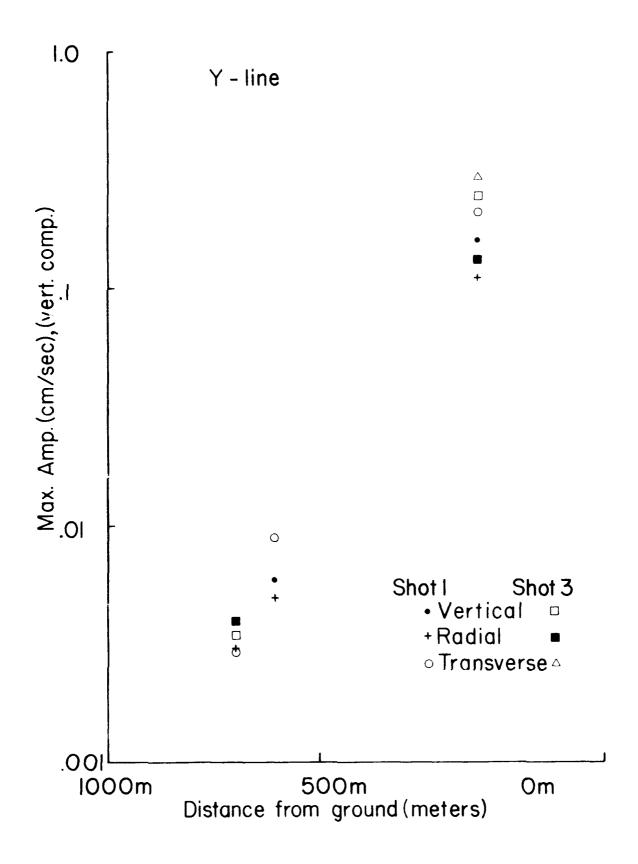


Figure 7. Maximum Amplitue vs Distance - Y-Line

A particle motion plot of FDY3 shot 1 is included in the next figure. Figures 8a, b, c, d, e, and f are particle-motion plots of selected stations. To understand these plots, consider Figure 8a, Station FDX1 (the closest station). The boxed area delineates the part of the waveform analyzed. The bottom 2 plots show that there was little transverse motion generated and that the rays came from the east-west (that is radial) direction. This pattern is repeated in Figure 8b, station FDX 2. As the box moves further into the waveform (Figure 8c), transverse motion begins to appear. In Figures 8d, e. and f. transverse motion is already apparent in the first motion portion of the record. Generally, the main cause of observed transverse motion from explosive sources is scattering of the waveforms by inhomogenieties along the ray path or reflections from a surface or boundary. Pressure waves, which are shown in idealized form in Figure 9, can be reflected off sides of hills, trees, and cloud layers. The shot day was clear, free of reflecting clouds. Figure 10 illustrates how the pressure wave from the explosions might reflect off the side of the Y-line hill and possibly off the trees defining the northsouth boundary of the wooded area. However, in this experiment there was no evidence of reflected acoustic energy on the pressure recordings. Figures 11 and 12 are record section plots using the vertical components for the X- and Y-lines, shots 1, 2, and 3. These plots are vertical seismogram components stacked in order of their distance from the source. The horizontal axis is "reduced time", which is computed by taking the observed travel time and subtracting the recording station's distance from the source divided by an estimate of the speed of the wave that is measured across the line.

Several points are apparent from these plots. First, station FDX5 has an obvious timing error. The change in frequency content of the waveforms Letween FDX3 and FDX4 and between FDY2 and FDY3 can be explained by the different instrumentation. (For close in stations the 200 sample per second recorders were paired with the 2 Hz natural period seismometers because the highest frequency signals were expected closer to the source.) For the X-line, a reduction velocity of 0.39 km/sec is shown. This lines up the waveforms of the first three stations, but FDX4 and FDX5 appear to be arriving late. This would suggest that either a slightly slower reduction velocity should be used or the propagation characteristics change at the farther stations, or the physics of the wave propagation at the farther stations is different. A velocity of 0.39 km/sec is quite slow for propagation of seismic waves in the crust. It seems likely that at least the close in stations were recording the direct air-blastcoupled ground wave. For the Y-line, a reduction velocity of 0.33 km/sec was used for the plot. However, the smaller number of stations and ambiguity in picking the onset of first motion on the 1 Hz (100 samples per second) instruments of FDY3 and FDY4, makes this number uncertain. To understand the physics involved, the geometry of the suspended shot must be considered. For the Y-line this geometry is complicated because of the hill. For the X-line a simple example is illustrated in Figure 13. In general, if the first motion energy is caused by the impact of the acoustic wave at the recording site, then the acoustic velocity is slower than the 0.39 km/sec apparent velocity used for the X-line.

In Figures 14 and 15, the waveforms recorded by the pressure transducers are displayed. Because of equipment limitations, one recorder was used for every 3 stations. The pressure sensors were attached with long cables. This setup had the drawback that if one recorder failed information at three stations would be lost. Because of the anticipated high frequency content of the signals, the higher sampling rate recorders were slated for this operation, but some last minute equipment problems necessitated substituting a 100 sample per sec instrument at the FDY sites. There were substantial equipment failures on these pressure sensor channels. The best data consisted of the

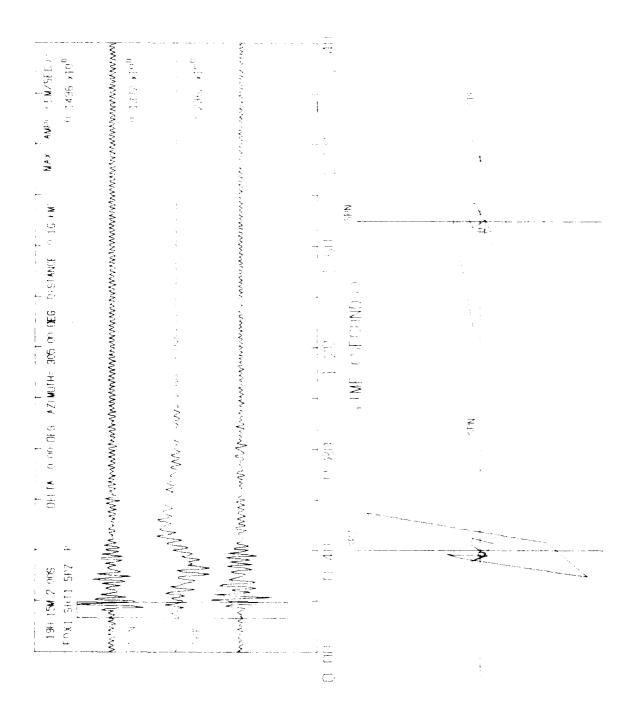
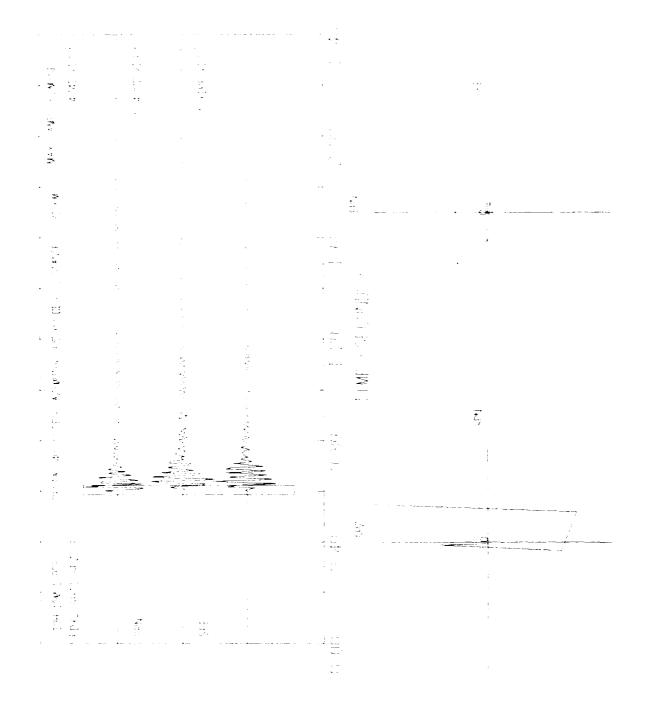


Figure 8. Particle Motion Plots



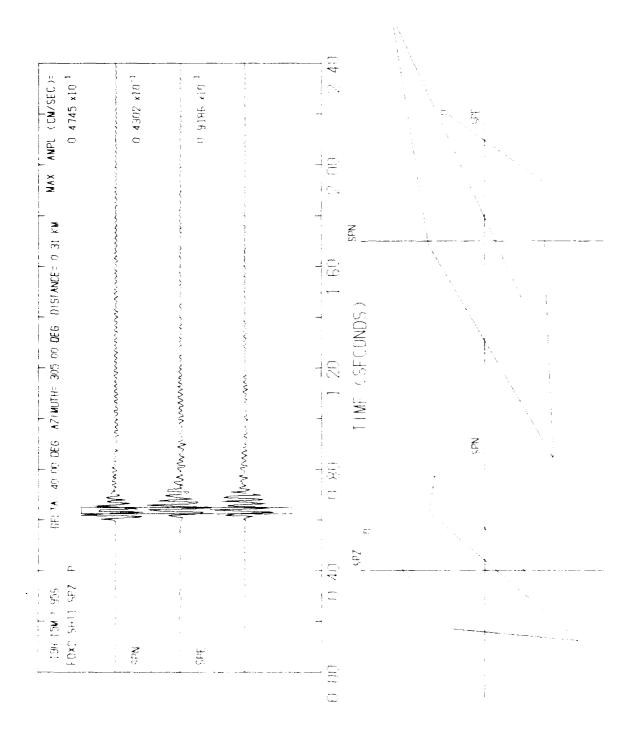


Figure 8. Particle Motion Plots (Cont'd)



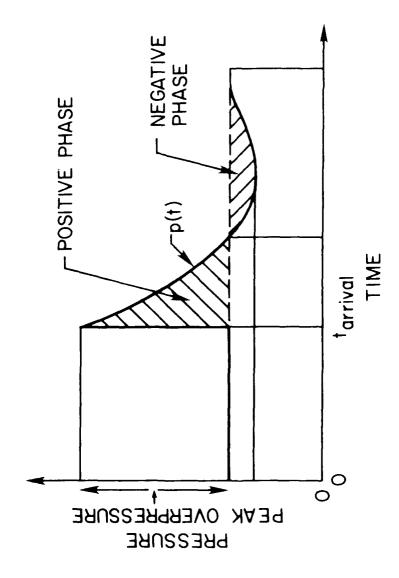


Figure 9. Ideal Blast Wave (Modified from Baker, 1973)⁷

7. Baker, W.E. (1973) Explosions in Air, University of Texas Press, Austin and London, pp 10-80.

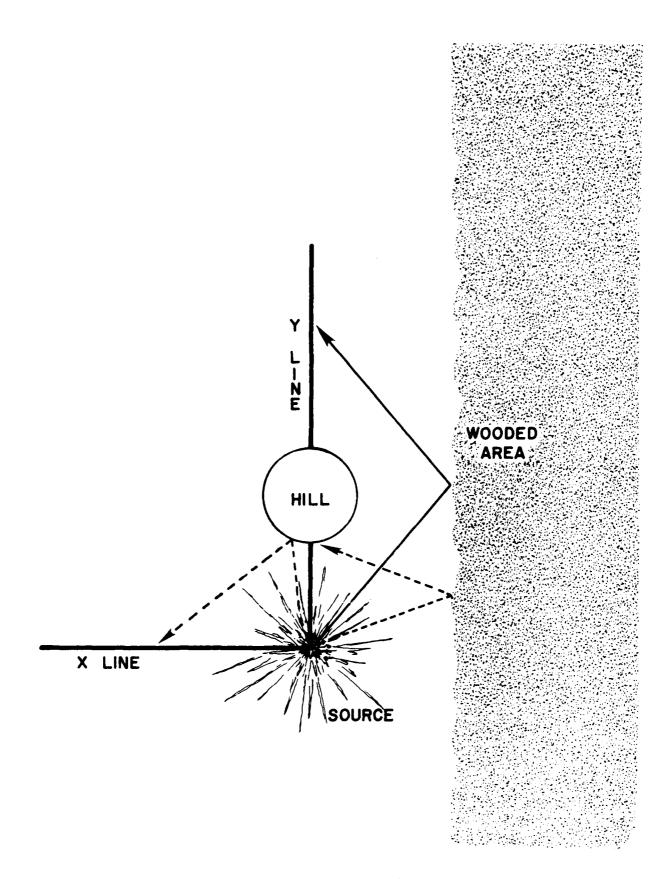


Figure 10. Possible Acoustic Wave Reflective Sources

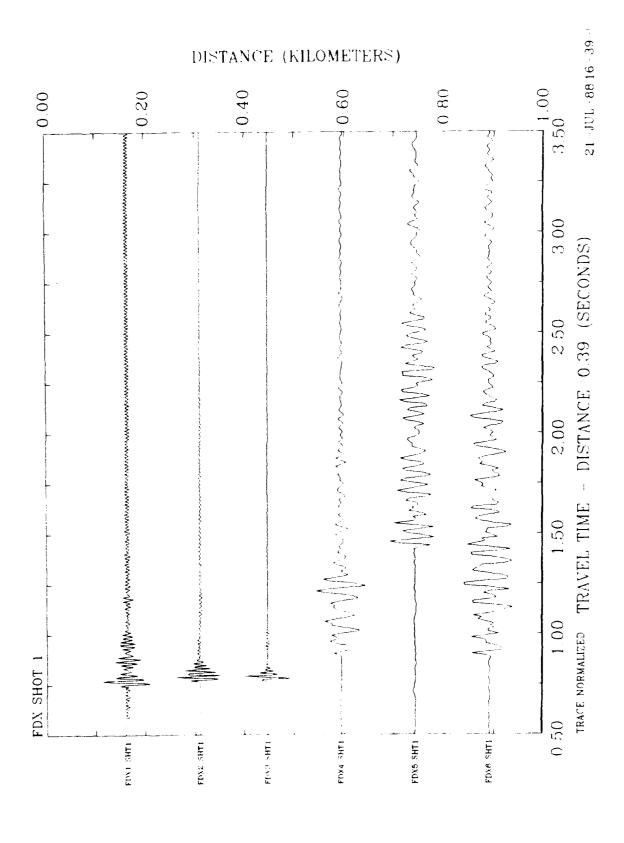


Figure 11. X-Line Record Section

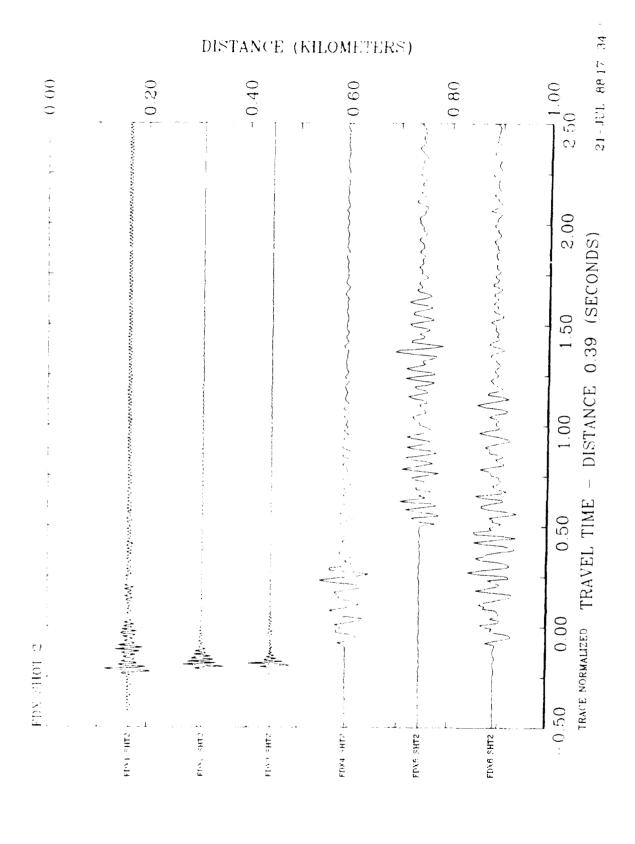


Figure 11. X-Line Record Section (Cont'd)

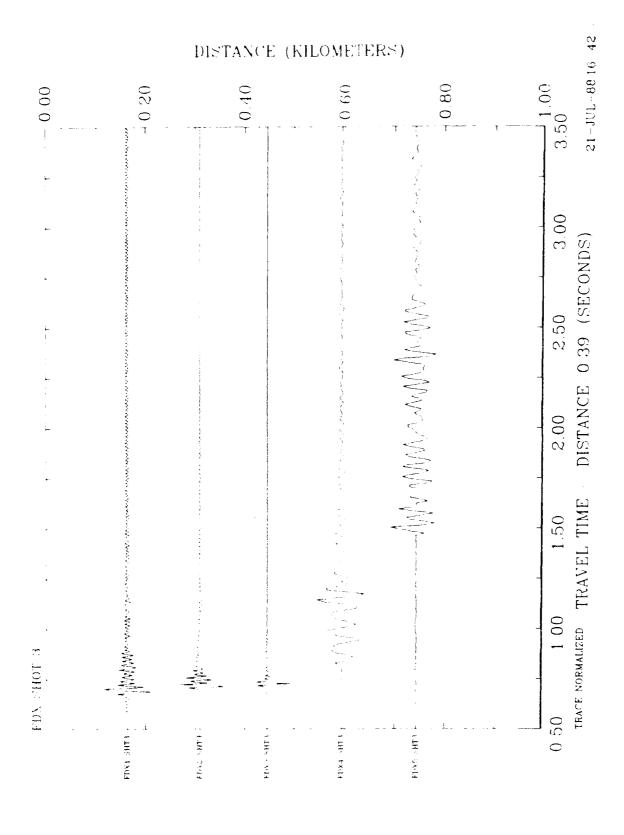


Figure 11. X-Line Record Section (Cont'd)

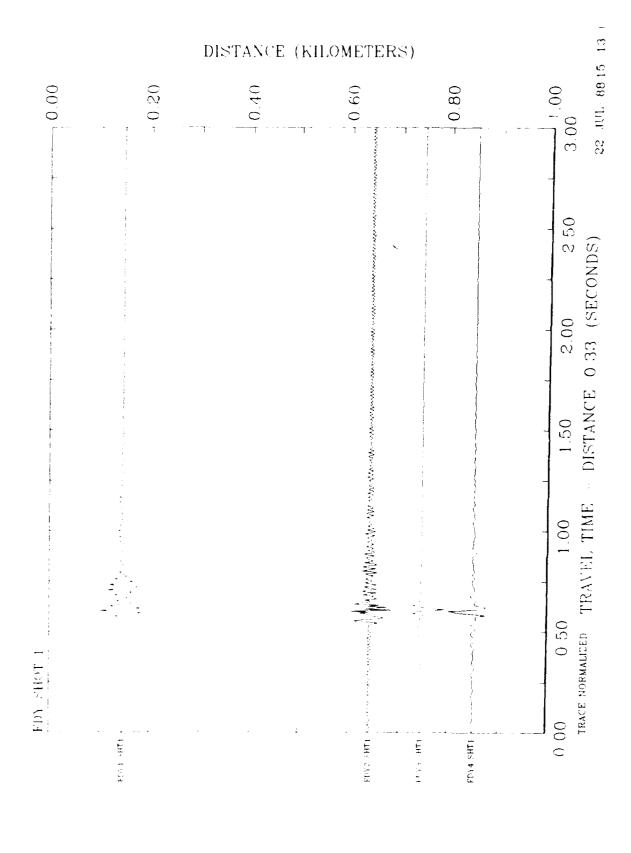


Figure 12. Y-Line Record Section

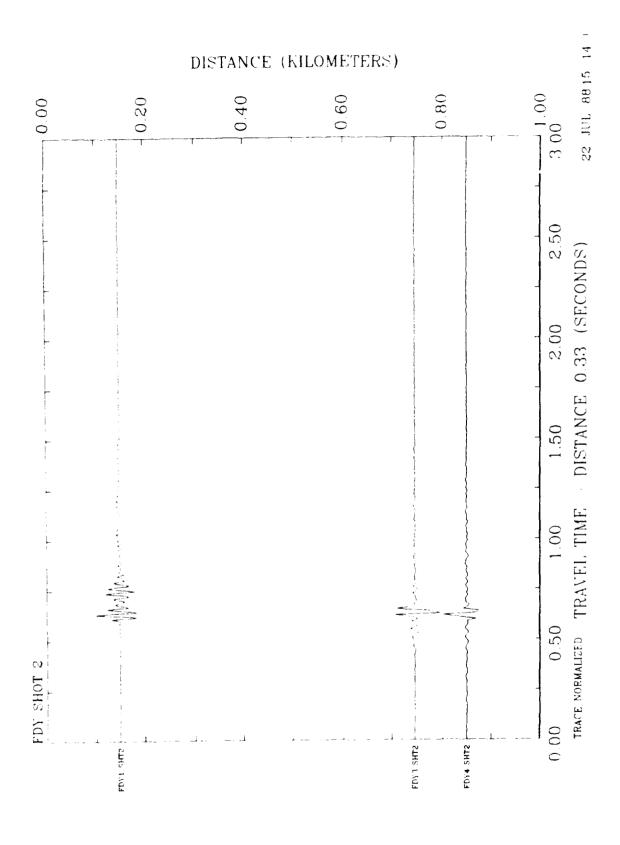


Figure 12. Y-Line Record Section (Cont'd)

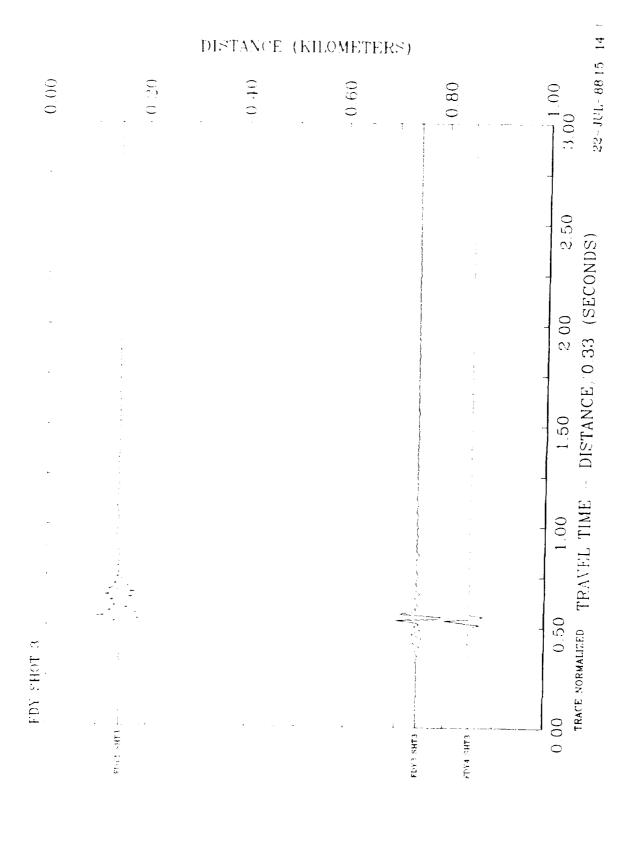


Figure 12. Y-Line Record Section (Cont'd)

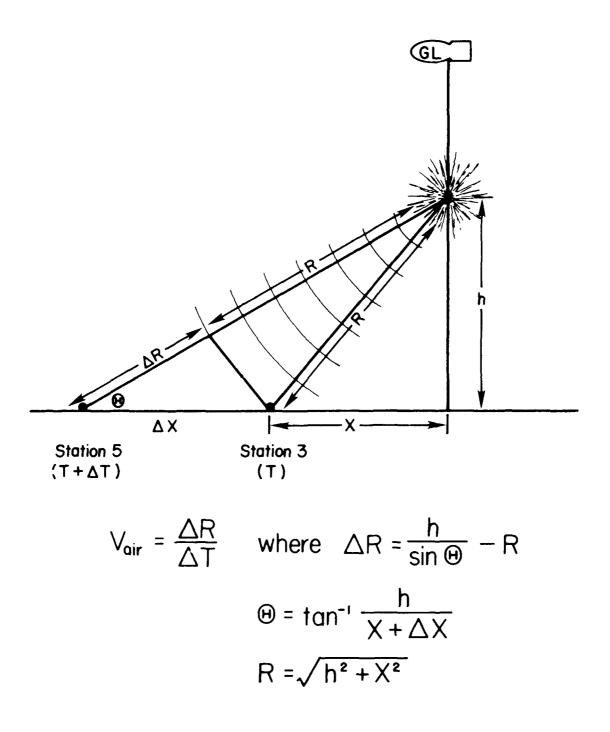


Figure 13. Air Wave Geometry

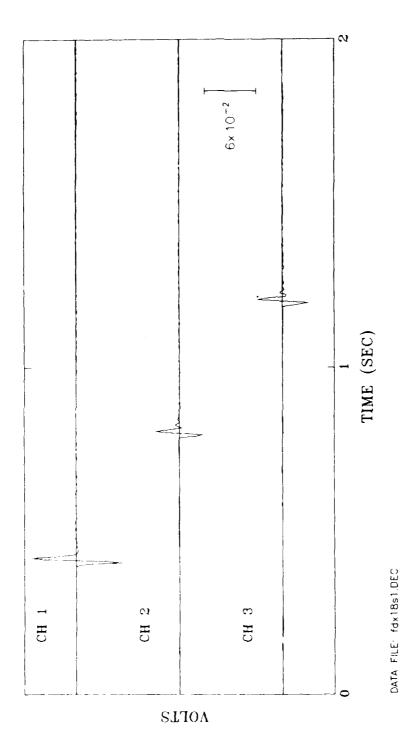


Figure 14. Pressure Sensor Data - X-Line

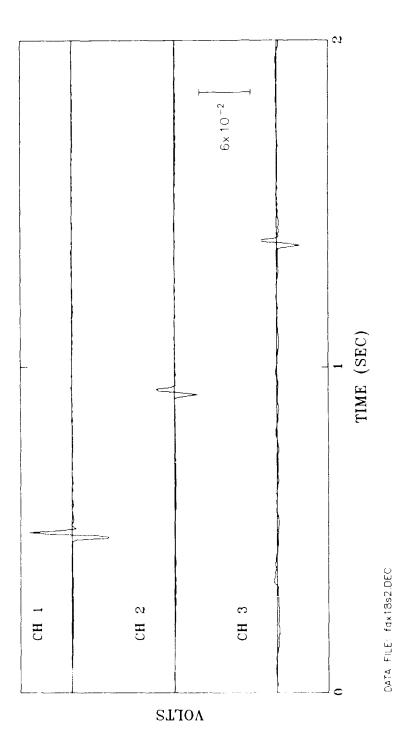


Figure 14. Pressure Sensor Data - X-Line (Cont'd)

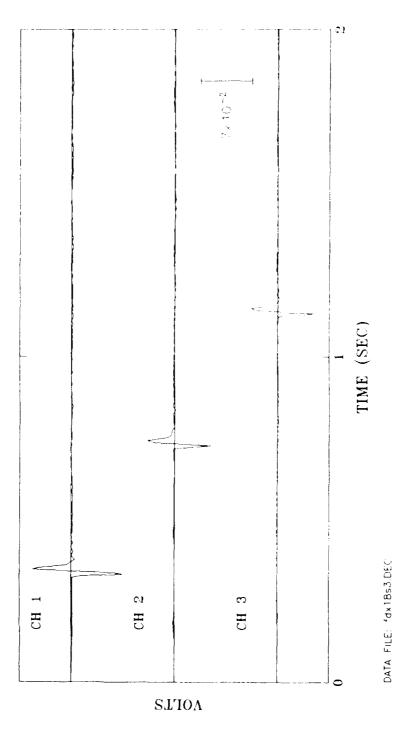


Figure 14. Pressure Sensor Data - X-Line (Cont'd)

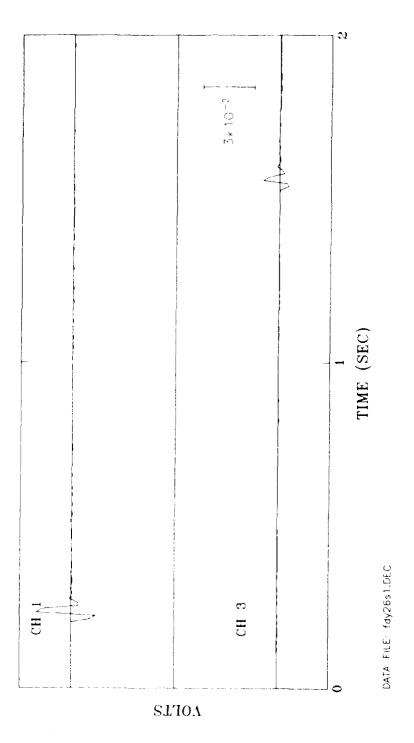


Figure 15. Pressure Sensor Data -Y-Line

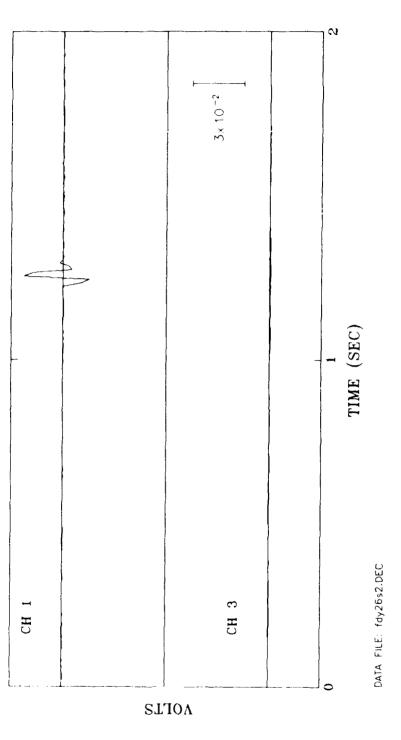


Figure 15. Pressure Sensor Data - Y-Line (Cont'd)

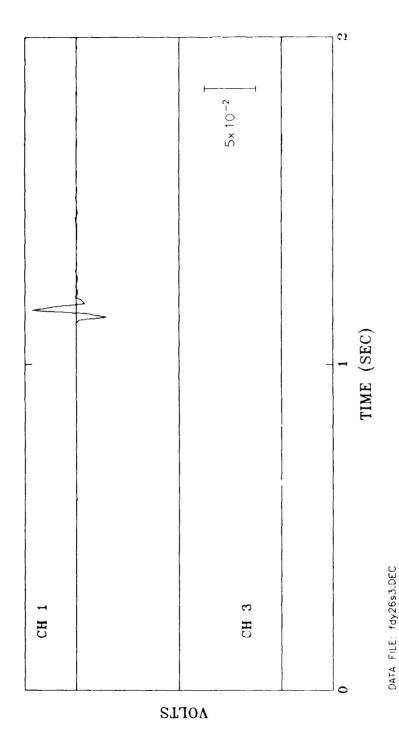


Figure 15. Pressure Sensor Data - Y-Line (Cont'd)

successful recording of all 3 shots at stations FDX4, FDX5, and FDX6. The results of seismo-acoustic calculations (ratios of peak seismic motion to acoustic amplitude) are shown in Table 2. The values

Shot 1 Shot 3 Station Distance (cm/sec/Pa) (cm/sec/P) (m) 1.7×10^{-6} 1.9×10^{-5} 593 FDX4 1.3×10^{-5} 1.4×10^{-5} FDX5 742 1.3×10^{-5} 891 FDX6 1.8×10^{-4} 147 3.0×10^{-4} FDY1 FDY2 643 4.0×10^{-4}

Table 2. Ratio of Peak Vertical Seismic Velocity to Peak Pressure

were in the range of 10^{-4} to 10^{-5} cm/sec/Pa. The higher values for the Y-line are probably a reflection of the higher sampling rate instruments at those stations. The values were consistent for shot 1 and shot 3 showing that for these soil conditions and this size source, admittance was independent of shot sizes.

Although the pressure profiles were plagued by equipment failures, some simple observations could be made. Some changes in travel time were probably the result of wind patterns and/or changes in temperature. (The temperature varied from 65 to 70° F during the shots; barometric pressure averaged 29.96 inches of mercury). The ratio of the maximum amplitudes of the third shot (10 lbs) to the maximum amplitudes of the first and second shots (5 lbs) ranged from 1.2 to 1.6. Coincidence of pressure wave and seismic wave arrival times shows that indeed for the X-line and the close in Y-line station, the measured seismic wave was directly induced by the pressure wave hitting the ground at the surface of the station.

However, the complexities of the coda indicate that some of the energy must result from seismic waves that were induced by the blast at other locations. For instance, an examination of a Y-line station behind the hill, FDY2 - shot 1 (Figure 5), shows a seismic precursor arriving at approximately 0.5 sec and a direct acoustically induced seismic wave at 0.65 sec. The precursor could be a seismic wave induced on the blast side of the hill that propagated through the hill at a higher velocity than the pressure wave (which must travel over the hill) and therefore arrived before the direct pressure wave. The hill was approximately 200 to 250 meters wide, so these arrivals are reasonable. Amplitude fall off with distance varied; more data points are needed.

Ratios of peak amplitudes between pairs of stations along the flat array, normalized to 100 meters distance, varied from 0.9 to 0.6. The ratio of peak seismic amplitude from the Y-line for stations FDY1 to FDY2, again normalized to 100 meters, was 0.3. FDY2 was behind the hill; however, because of the evaluation of the shot, the geometry was such that station FDY2 was at a closer straight line distance to the shot (see Figure 3).

5. DISCUSSION

From the seismic data, it is clear that the hill sheltered the Y-line stations, and that doubling the amount of explosives increased the peak amplitudes by a factor of 1.1 to 1.3. Considerable transverse motion is apparent on the records. Particle motion plots showed that the transverse motion occurred earlier in the record as the recording distance increased. Reduced time plots show a propagation speed of approximately 0.35 km/sec. For the acoustic data, seismo-acoustic peak amplitude ratios in the range of 10^{-4} to 10^{-5} cm/sec/Pa were observed. Ratios of peak pressure for the larger shot (10 lb TNT) to the smaller shots (5 lb TNT) were in the range of 1.2 to 1.4. The experiment was repeatable: waveforms from shot 1 and shot 2 are nearly identical.

This experiment was performed as a feasibility study for determining the effect of topography, terrain, and vegetation on pressure and pressure-induced seismic signals for low altitude atmospheric explosions. Because of the frequencies involved, follow up tests should employ systems capable of higher sampling rates so that a major percentage of the radiated energy will be recorded. The Turner Drop Zone itself is a very good area to use for a comparison of the effect of snow cover on the signals if the experiment is repeated in the winter. Even with the equipment limitation, it was possible to determine that some of the seismic waves were directly induced by the acoustic wave, and amplitude ratios for changing source strength were obtained as well as some attenuation data. This experiment can easily be repeated in areas of differing vegetation, ground cover, and topography. It is recommended that in follow-up experiments the position of the shot be varied so that the geometry will be changed enough to discriminate the various seismic paths induced by the blasts. GL has continued its program of seismo-acoustic research since the Ft Devens experiment with recordings of B-1 aircraft flyovers⁸ and measurements of the ambient seismo-acoustic environment.⁹

^{8.} Crowly, F.A. and Blaney, J.I. (1987) Surface Disturbances Produced by Low Level, Subsonic B-1 Aircraft, AFGL-TR-87-0325, ADA192257.

^{9.} Battis, J.C. and Center, C. (1989) Temporal Attributes of the Ambient Seismo-Acoustic Environment: La Junta, CO, AFGL-TR-89-0080, ADA219369.

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- 3. United States Army Forces Command, 584th Engineer Company (Cartographic), 30th Engineer Battalion (Topographic) (Army), Fort Belvoir, VA 22060, (Feb, 1986) Fort Devens Special Map, edition 2-DMATC, Series V8145.
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- 5. Nicholls, M.R., Johnson, C.F., and Duval, W.I. (1971) Blasting Vibrations and Their Effects on Structures, Bulletin 656, US Dept. of the Interior, Bureau of Mines, pp. 65-66.
- 6. Battis, J.C. (1985) Vibro-Acoustic Forecasts for STS Launches at V23, Vandenberg AFB: Results Summary and the Payload Preparation Room, AFGL-TR-85-0013, ADA162192.
- 7. Baker, W.E. (1973) Explosions in Air, University of Texas Press, Austin and London, pp 10-80.
- 8. Crowly, F.A. and Blaney, J.I. (1987) Surface Disturbances Produced by Low Level, Subsonic B-1 Aircraft, AFGL-TR-87-0325, ADA192257.
- 9. Battis, J.C. and Center, C. (1989) Temporal Attributes of the Ambient Seismo-Acoustic Environment: La Junta, CO, AFGL-TR-89-0080, ADA219369.

Appendix A: Equipment Parameters

Table A-1. Clock Correction Data
Time Slew Relative to Recorder 301 at End of Shot Day

Recorder Number	Sample Rate (s/sec)	Slew*	Time Correction (msec)
284	200	550	-83
283	200	580	-50
269	200	320	-466
301	200	000	0
270	200	(turned off)	
300	200	583	-28
299	200	544	-93
281	100	007	+23
277	100	010	+33
322	100	292	-27
320	100	104	+346
280	100	004	+13
282	100	273	-90
312	100	005	+16

^{*} Slew is relative to 300 for 100 sps or 600 for 200 sps recorders.

Table A-2. Pressure Sensor Calibration Data Scale Factors*

Sensor Number	Range (psi)	+0.5 psi (V)	0.0 psi (V)	-0.5 psi (V)
1	±0.5	+0.375	+0.009	-0.389
2	±0.5	+0.684	-0.001	-0.656
3	±0.5	+0.499	+0.001	-0.502
4	±0.5	+0.567	-0.022	-0.511
5	±1.5	+0.213	+0.000	-0.214
6	±1.5	+0.258	-0.000	-0.252
7	±1.5	+0.219	-0.000	-0.218
8	±1.5	+0.209	+0.001	-0.212

^{*} Using (± 6.92) 13.84 inches of H₂O = 0.5 psi

Table A-3. Recorder, Seismometer and Pressure Sensor List

		Se	ismic				Pressu	ıre	
Station	Recorder	Sampling	Seisn	nomete	rs	Recorder	Sampling	РТ	
Location		Rate	Channel 1	2	3		Rate		
X1	270	200	8296Z	NS	EW			0.5 PSI	#6
X2	283	200	7493Z	NS	EW	284	200	0.125 PSI	#7
Х3	299	200	7679Z	NS	EW			0.125 PSI	#5
X4	312	100	540Rad	Т	556Z			0.125 PSI	#66
X5	320	100	546N	543E	551Z	300	300	0.125 PSI	#
Х6	322	100	548N	542E	553Z			0.125 PSI	#
Yl	269	200	7680Z	NS	EW			0.5 PSI	#
Y2	301	200	9261Z	NS	EW	281	100	0.5 PSI	#
Y3	280	100	549N	545E	555Z			0.5 PSI	#
Y4	277	100	547N	539E	558Z				

Appendix B: X- and Y-Line Properties

"" INDEX OF ARRAY SEISMOGRAMS ""

					мо	DY	YEAR	HR	MN	SECOND	DISTANCE	DEI.TA	AZIMUTH
1	FDX1	SHT1	SPZ	P	5	5	1986	19	15	2.000	0.16	0.00	305 00
9	FDX1	SHT1	SPN	P.	5	5	1986	19	15	2.000	0.16	0.00	305 00
	FDX1	SHT1	SFE	P	5	5	1986	19	15	2.000	0.16	0.00	305 00
4	FDM1	SHT2	SFZ	P	5	5	1986	19	39	49.000	0.16	0.00	305 00
ē	FDXI	SHTS	SPN	F	5	5	1986	19	39	49.000	0.16	0.00	305 00
-6	FDX1	SHTS	SPE	P	5	5	1986	19	39	49.000	0.16	0.00	305.00
7	FDX1	SHT3	SPZ	P	5	5	1986	20	3	39.000	0.16	0.00	305.00
8	FDX1	SHT3	SPN	P	5	5	1986	20	3	39.000	0.16	0.00	305.00
9	FDX1	SHTT	SPE	P	5	5	1986	20	3	39.000	0.16	0.00	305.00
10	FDX2	SHT1	SPZ	P	5	5	1986	19	15	2 000	0.31	40.00	305.00
11	FDX2	SHTl	SPN	P	5	5	1986	19	15	2.000	0.31	40.00	305 00
12		SHTl	SPE	P	5	5	1986	19	15	2.000	0.31	40.00	305.00
13		SHTS	SPZ	P	5	5	1986	19	39	49.000	0.31	40.00	305.00
14	EDX S	SHTS	SPN	P	5	5	1986	19	39	49.000	0.31	40.00	305.00
15		SHIS	SPE	F	5	5	1986	19	39	49 000	0.31	40.00	305 00
16		SHT3	SPZ	P	5		1986	80	3	39 000	0.31	40.00	305.00
17	FDXS	SHT3	SPN	P	5	5	1986	20	3	39.000	C.31	40.00	305 00
18		SHT3	SPE	P	5	5	1986	80	3	38 000	0.31	40.00	305.00
20	FDX3	SHTI	SPZ	P	5	5	1986	19	15	2.000	0.44	20.00	280 00
21	FDX3	SHT1	SPN	P P	5	5	1986	19	15	2.000	0.44	20.00	280.00
55	FDX3	SHT2	SPE	-	5	5	1986	19	15	2.000	0.44	20.00	260.00
23	FDX3		SPZ	P	5	5	1986	19	39	49.000	0.44	20.00	280.00
24	FDX3	SHT2	SPN	P	5	5	1986	19	39	49.000	0.44	20.00	280.00
25	FDX3	SHIZ	SPE	P P	5	5	1986	19	39	49.000	0.44	20.00	280.00
26	FDX3	SHT3	SPN	P P	5 5	5	1986	20	3	39.000	0.44	20.00	280 00
27	FDX3	SHT3	SPE	P P	5	5 5	1986 1986	20	3 3	39.000	0.44	20.00	280.00
28	FDX4		RADL		5	5	1986	19	15	39.000	0.44	20.00	280.00
ဥပ	FDX4	SHT1	TRNS	P	5	5	1986	19	15	2.000	0.59	80.00	295.00
40	FDE4	SHT1	SPZ	P	5	5	1986	19	15	2.000	0.59	80.00	295.00
7 -		SHTS	RADL	-	5	5	1986	19	39	2.000 49 .000	0.59	80.00	295.00
., .		HIP	TRNS	P	5	5	1986	19	39	49.000	0.59 0.59	80 00	295 00
د ک		SHIP	SPZ	P	5	5	1986	19	39	49.000	0.59	80.00 80.00	295 00 295 00
34	FDX4	SHT3	RADL	P	5	5	1986	50	3	39 000	0.59	80.00	295.00
3.2	FDM4	SHT3	TRNS	P	5	5	1986	20	3	39.000	0.59	80.00	295.00
36	FDX4	SHT3	SPZ	P	5	5	1986	20	3	39.000	0.59	80.00	295.00
3 1.	FINE	SHT1	SPN	P	5	5	1986	19	15	2.000	0.74	30.00	300 00
?8	$FL \times^{\alpha}$	SHTI	SPE	P	5	5	1986	19	15	2.000	0.74	30.00	300.00
.≯ C1	$FT(X^{\mathfrak{m}})$	SHT1	SPZ	P	5	5	1986	19	15	2.000	0.74	30.00	300.00
40	FDX5	SHT2	SPN	P	5	5	1986	19	39	49.000	0.74	30.00	300.00
41	FDX5	SHTS	SPE	P	5	5	1986	19	39	49.000	0.74	30.00	300.00
42	FDX5	SHT2	SPZ	P	5	5	1986	19	39	49.000	0.74	30.00	300.00
4.7	FDX5	SHTE	SPN	P	5	5	1986	80	3	39.000	0.74	30.00	300.00
44	FDX5	SHT3	SPE	P	5	5	1986	20	3	39.000	0.74	30.00	300 00
4.0	FDX5	SHT3	SPZ	P	5	5	1986	50	3	39.000	0.74	30.00	300 00
46	FDX6	SHTI	SPN	F	5	5	1986	19	15	2.000	0.89	0.00	305 00
4.7		SHTI	SFE	P	5	5	1986	19	15	8.000	0.89	0 00	305 00
4.8	FDX6		SFZ	F	5	5	1986	19	15	2.000	0.89	0 00	305 00
49	FDX6	SHIR	SPN	P	5	5	1986	19	39	49.000	0.89	0.00	305 00
50	FDX6	SHTS	SPE	P	5	5	1986	19	39	49.000	0.89	0.00	305 00
5.1	FDX5	SHTS	SFZ	P	5	5	1986	19	39	49.000	0.89	0 00	305 00

Table B1. Seismometer Constants (X-Line Properties)

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m)					c-4				. 4	Ŋ.	Ą.	N	51				<u>.</u>	y e	9		٠. :		۲	~	FDX3	۲	7	+	+	.,	, 4	,	,	.,,	1		L	E - 1				ι		ာ	6	,

Table B1. Seismometer Constants (X-Line Properties) (Cont'd)

SEISHOGRAM	STARTING TIME	TIME CORRECTION (SECONDS)	SEISMOMETER ORIENTATION (DEGREES)	SEISMOMETER PENDULUM SENSITIVITY PERIOD (VOLIS M SEC) (SECONDS)		DAMPING RATIO	SERIAL NUMBER	CALIBRATION DATE	NO
FDX6 SHI2 SPN	5 MAY 1986 125 19 39 49 000	-0.0270 SLEW	15.00	1152.5781	1.031	0.321	104548	O JAN	0
FDX6 SHT2 SPE	5 MAY 1986 125 19 39 49.000 -0.0270 SLEW	-0.0270 SLEW	75.00	1100.4688	1.020	0.321	104542 0 JAN	OJAN	0
FDX6 SHT2 SPZ	5 MAY 1986 125 19 39 49.000	0.0270 SLEW		1183.4646	1.064	0.329	104553	OJAN	0

Table B-2. DCS-302 Recorder Constants (X-Line Properties)

		FREQUENCY (HERIZ)	GAIN 1	GAIN 2	GAIN 3	GAIN 4	NUMBER
INS	SPZ	30 00	4414	4882	0826	0244	870
LHS 1	NASI		4414	488	0	244	270
		0 '	4414	4882	0976	0.844	270
TDX1 SHI		<u>,</u>	44.4	4882	.0976	0.24	270
SHI			4434	488	.0976	0	270
SHI		0	4414	4882	.0976	0.02441	270
SHI	3 SPZ		4414	4882	.0976	0.02441	270
SHI			4414	4882	9460	0.02441	270
SHI			4414	4882	.0978	.0244	270
THS		30 00	4414	488	087	œ	288
FUXE SHI			* 6 2 4	4.88	0876	.0244	283
THO			4414	4882	.0978	0	283
IHS	SPZ		4414	4882	0876	0.02441	283
SHI			4414	4882	9460	0.02441	283
SHI		30.00	4414	488	0876	0.02441	283
FDX2 SHI			4414	4882	.0978	0	283
H	N C L	30.00	441	4882	0	0244	283
1 5			4414	488	.0978	.0244	283
1 1 1			44]4	4882	0	0.02441	588
SHI			441	4882	0976	0.02441	868
IHS			441	4882	0976	0.8	888
1 1 0			4	4882	0.0976	0	599
THE STATE	200		44	4882	0976	0244	599
1 H	1 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	00.00	919	00 0	0.09788	244	588
1 1			1 th 1 th 1 th 2	200	.0978	0	868
H			# *	200	<u>٠</u>	8	568
1 1	ייני דרקט דרקט			2000	92.60	0244	899
H		00.0	0.00	5 0	0087	0024	318
1 1 1				Ö	0088	.0084	318
1 T			8 0)		0	312
SHT				5 6	0.00974	0024	312
THS 4	27.5		. 6	5 6		٠	318
THS #		00.08	4	. (318
4 SHT			. 0	5 6		# X 0 0 0	315
4 SHI	SPZ	C	0447	5 6		4300.	318
SHI		C	3 0			0083	315
SHI		, C	3 0	5	200	4200	350
S		0	9 0	Ċ	08800		320
5 SHT		¢				4800.	350
5 SHT	SPE			. 0		0.00244	320
5 SHT) C	24.0	5 6	900	4800.	380
5 SHT		. c		5 6	8600.	.0024	350
5 SHT		, ,	0 P P C		8600.	.0024	380
SSHT) C	10 14 C	.	6600	0.24	320
S SHT) (* u	5 0	.0098	.0024	320
S				5 6	800.	.0024	3,22
SHT) (0 0	* •	8600	003	322
•		?	. K	4	0.00981	0.00242	322

Table B-2. DCS-302 Recorder Constants (X-Line Properties)(Cont'd)

SERIAL Number	3 8 8 8	322
COUNT) GAIN 4	0.00245	0.00242
RECORDER GAINS (MILLIVOLIS COUNT) N 1 GAIN 2 GAIN 3 GAIN 4	0.00970	0.00981
RECORDER GAINS (0.04907	0.04973
RECOR GAIN 1	0.25326	0.25562
CUT OFF FREQUENCY (HERIZ)	30 00	30.00
SEISHOGRAH	FDX6 SHI2 SPN	FDX6 SHT2 SPZ

INDEX OF ARRAY SEISMOGRAMS ' ' :

					мо	DY	YEAR	HR	MN	SECOND	DISTANCE	DELTA	AZIMUTH
1	FDY1	SHT1	SPZ	P	5	5	1986	19	15	2.000	0 15	0 00	c 00
ē.	FDY1	SHT1	SPN	P	5	5	1986	19	15	2.000	0.15	0.00	0.00
3	FDY1	SHT1	SPE	P	5	5	198€	19	15	2.000	0.15	0.00	0 00
4	FDY1	SHTS	SPZ	P	5	5	1986	19	39	49.000	0.15	0.00	0.00
5	FDY1	SHT2	SPN	P	5	5	1986	19	39	49.000	0.15	0.00	0.00
6	FDY1	SHT2	SPE	P	5	5	1986	19	39	49.000	0.15	0.00	0.00
7	FDY1	SHT3	SPZ	P	5	5	1986	20	3	39.000	0.15	0.00	0.00
8	FDY1	SHI3	SPN	P	5	5	1986	20	3	39.000	0.15	0.00	0.00
8	FDY1	SHT3	SPE	P	5	5	1986	20	3	39.000	0.15	0.00	0.00
10	FDY2	SHT1	SPZ	P	5	5	1986	19	15	2.000	0.64	0.00	0.00
11	FDY2	SHT1	SPN	P	5	5	1986	19	15	2.000	0.64	0.00	0.00
12	FDY2	SHT1	SPE	P	5	5	1986	19	15	2.000	0.64	0.00	0 00
13	FDY3	SHIL	SPN	P	5	5	1986	19	15	2.000	0.74	0.00	0.00
14	FDY3	SHTl	SPE	P	5	5	1986	19	15	2.000	0.74	0.00	0.00
15	FDY3	SHTl	SPZ	P	5	5	1986	19	15	2.000		0.00	0.00
16	FDY3	SHIS	SPN	P	5	5	1986	19	39			0.00	0.00
17	FDY3	SHT2	SPE	P	5	5	1986	19	39	49.000		0.00	0.00
18	FDY3	SHIS	SPZ	P	5	5	1986	19	39	49.000		0.00	0.00
19	FDY3	SHT3	SPN	P	5	5	1986	-	3	39.000		0.00	0.00
80	FDY3	SHT3	SPE	P	5	5	1986		3	39 000		0 00	0.00
21	FDY3	SHT3	SPZ	P	5		1986		3			0.00	0.00
22	FDY4	SHTI	SPN	P	5		1986		15	2.000		0.00	15 00
23	FDY4		SPE	P	5		1986		15	2.000		0.00	15 00
24	FDY4		SPZ	P	5		1986		15	2.000		0.00	15 00
25	FDY4	SHTS	SPN	P	5		1986		39			0.00	15 00
26	FDY4		SPE	P	5			-	39			0.00	15.00
5~	FDY4	SHTP	SFZ	P	5		1986		39			0 00	15.00
· 8	FDY4	c H I a	CEN	P	5		1986	-	3			0 00	0.00
- "				E.	5				3			0 00	
3	FDY4	SHIB	SPZ	F	5	5	1986	20	3	39.000	0.85	0.00	0.00

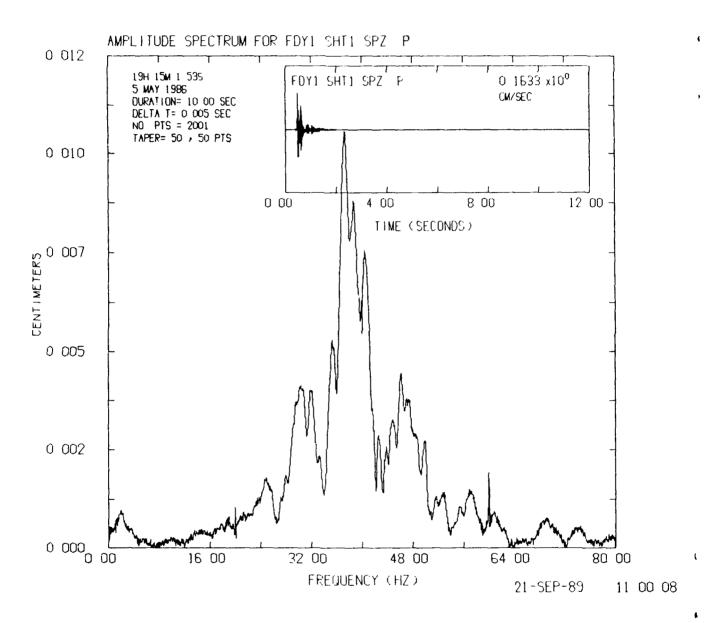
Table B-3. Seismometer Constants (Y-Line Properties)

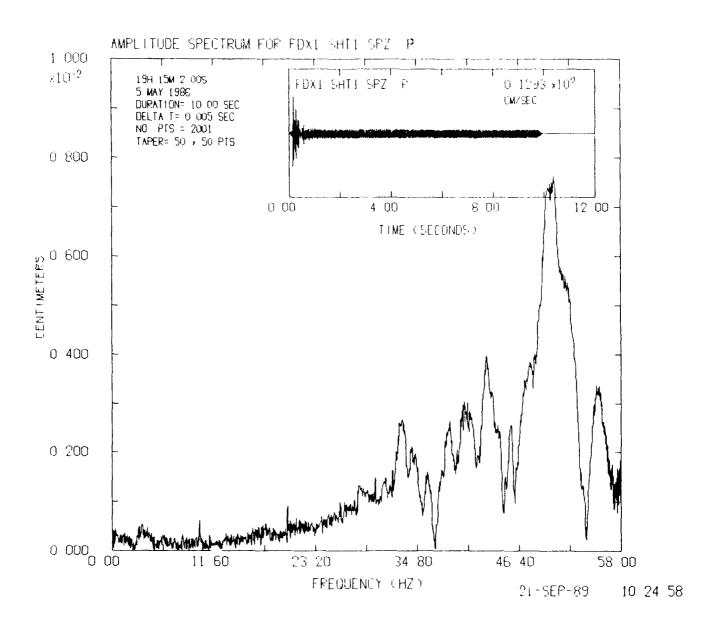
SEISMOGRAM	OGRAM		STARTING		TIME		CORRECTION (SECONDS)	ORIENTATION (DEGREES)	SENSITIVITY (VOLTS M SEC)	PERIOD (SECONDS)	RATIO	NUMBER	DATE	•
FLY1 SHT	T1 SFZ	S MAY	1986		-	ļ	0 4660 SLEW		200.0000	0 500	0.400	7680	O JAN	0
FDY1 SHT1	S	5 MAY			7	O)	4660	345 00	200.0000	0 500	0.400	7680	O JAN	0
		5 MAY	œ		7	8		75.00	800.0000	0.500	0.400	7680	O JAN	0
_	S	5 MAY	œ	125	ы	9	0.4860 SLEW		200.0000	0.500	0.400	7680	OJAN	0
FDY1 SHT2	SP	5 MAY	1986	125	19 3	9 49.000	0.4660 SLEW	15.00	200.0000	0.500	0.400	7680	O JAN	0
_	I2 SPE	S MAY	1986	S.	9	4	0.4660 SLEW	75.00	200.0000	0.500	0.400	7680	O JAN	0
	T3 SPZ	5 MAY	1986	125		n	0.4660 SLEW		200 0000	0.500	0.400	7680	OJAN	0
		5 MAY	1986	S	0	E)	0.4660 SLEW	15.00	200.0000		0.400	7680	O JAN	O
FDY1 SHI3		5 MAY	1986	ខ	0	n	0.4660 SLEW	75.00	800.0000		0.400	7680	O JAN	O
	II SPZ		œ	S.	9		0000		200 0000		0.400	9261	OJAN	O
			æ	185	-		0000	-15.00	200.0000	0.500	0.400	9261	O JAN	O
	II SPE	5 MAY	1986	125	٦		0 0000 SLEW	75.00	200.0000	0.500	0 400	9261	O JAN	_
FDY3 SHT	_	5 MAY	1986	125	~	S.	0 0130 SLEW	-15.00	1152 5391	1.042	0 337	104549	O JAN	_
			1986	125	_	ហ	0 0130 SLEW	75.00	1109.1016	1.042	0 314	104545	O JAN	
FDY3 SHT1	T1 SPZ		œ	125	~	ស	0.0130 SLEW	1		1.053	0.325	104555	O JAN	O
			98	œ	B	9.49	0130	- 15 .00		1.042	0.337	104549	OJAN	Č
			98	œ	9	648	0130	75.00		1.042		104545	OJAN	
			98	83	9	6.6		f		1.053	•	104555	OJAN	0
n	S		1986	ຜ	0	n	0.0130 SLEW	- 15.00	1152.5391	1.042	0.337	104549	OJAN	
(*)	3 SP		1986	125		B	0.0130 SLEW	75.00	1109.1016	1.042	0.314	104545	O JAN	
	SP		98	52	0	t)			1123.5039	1.053	0.325	104555	OJAN	Č
			æ	82	9	αį		15.00	1103.1250	1.042	0.306	104547		•
4	S		Φ	Ω	9	5 10	.0330	75.00		1.020		104539	NVP O	Ŭ
-41	S		œ	ED.	9	n G	.0330			1.042		104558	OJAN	
FDY4 SHT2			œ.	ın	n	9 49.	.0330	-15.00		1.042	0.306	104547	OJAN	Ŭ
-44	SP		Ø	125	9	9.49	.0330	25.00	1123.0859	1.020	0.310	104539	OJAN	_
·	SE		œ	വ	9	9.49	.0330	1	1218.7402	1.042	0.346	104558		Ŭ
ν: •	SP		œ	വ	0	n	0.0330 SLEW	-15.00	1103.1250	1.042		104547		_
4. ب	S		1986	125		39	0330	75.00	1123 0859	1.020	0.310	104539		_
FDY4 HT	S. P.	S MAY	α	S	0	n	0.0330 SLEW		1218,7402	1.042	0.346	104558	OJAN	_

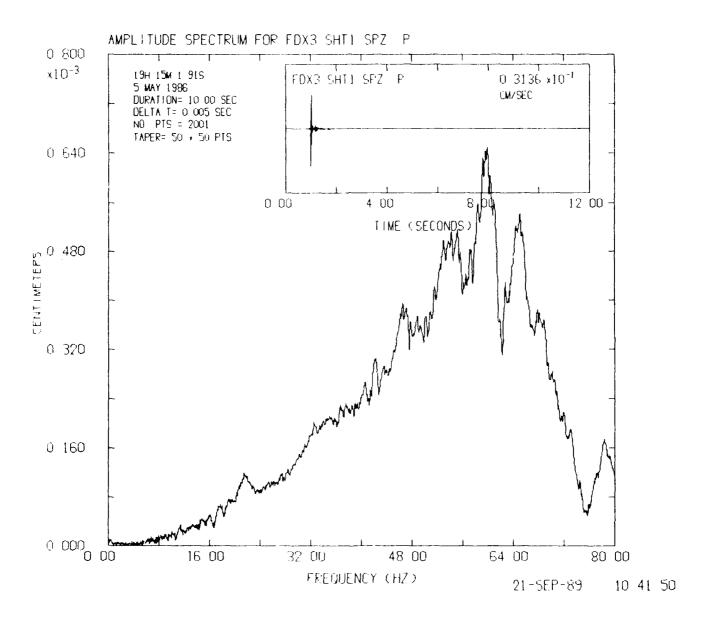
Table B-4. DCS-302 Recorder Constants (Y-Line Properties)

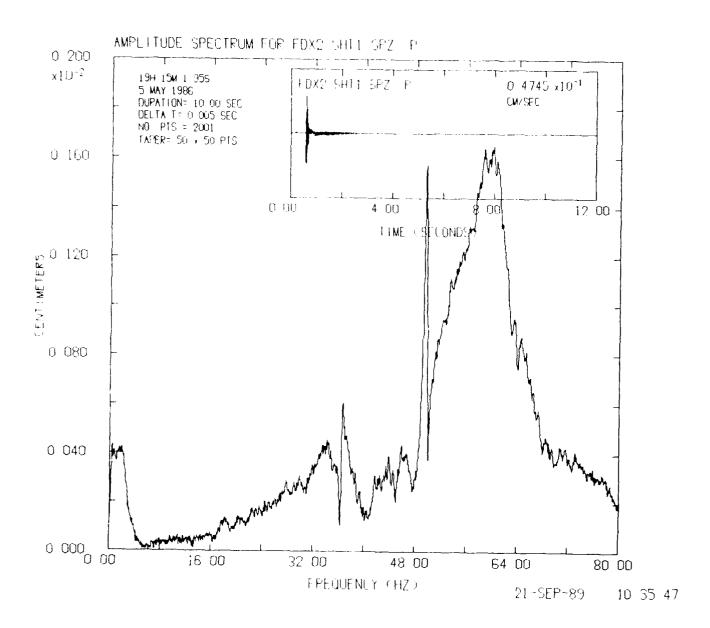
9EISMOGRAM	FREQUENCY	GAIN 1	N 1 GAIN 2	GAIN 3	COUNT) GAIN 4	SERIAL
HT1 SPZ	30 00	2 44141	0.48828	0 09766	0 02441	080
HT1 SPN	30 00	4	-	0.00	400	000
HII SPE		2 44141	0 48828	0.09766		0 0
	30 00	4.4	488	9260	0.20	0 0
SSF	30 00	2 44141	4882	9260	0.24	000
HI2 SPE	30.00	2 44141	0.48828	0976	0.24	269
	30 00	2 44141	0.48828	0.09766	0.5	269
က လ	30 00	2.44141	0.48828	0.39786	0.02441	869
۲,	30 00	2.44141	0.48828		0.02441	269
	30 OO	2 44141	48	0.09766	0.02441	301
	30 00	2.44141			024	301
		8.44141			.024	301
	30 00 30 00	Q			.0023	980
		2453				280
		8.45		0.00977	0 00235	280
SPN		245	0.04947	0.00978	0023	280
		2453		0.00980	0023	280
	-	45				280
Z. 1.		3.45			0.00235	280
		2453			.0023	280
		6 4 10			0083	280
SHII SFN	-	838		٠.	0.00238	277
		2385	-	0.00964	0.00246	277
	30 OC	. 238	0.	0.00971	0.00246	277
27. E		2389		0.00974	0.00238	277
		. 2385	0	0.00964	0.00246	277
	30 00	38	.	0.00971	.0024	277
S	_	2389	0.05039	0.00974		277
S	30 70	. 2385	0.	0.00964	0.00246	277
ZES	00 00 00	0.23872	0.04906	[2000 0	0.000	1

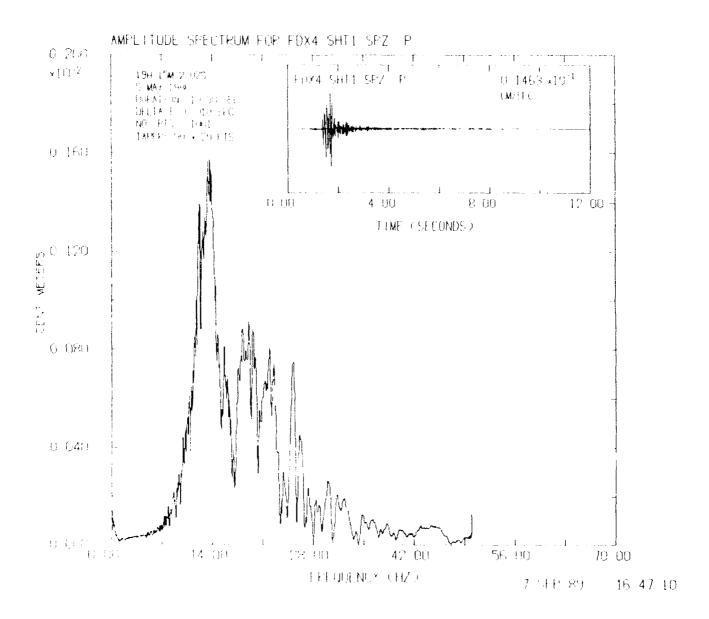
Appendix C: Amplitude Spectra for Selected Seismograms











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